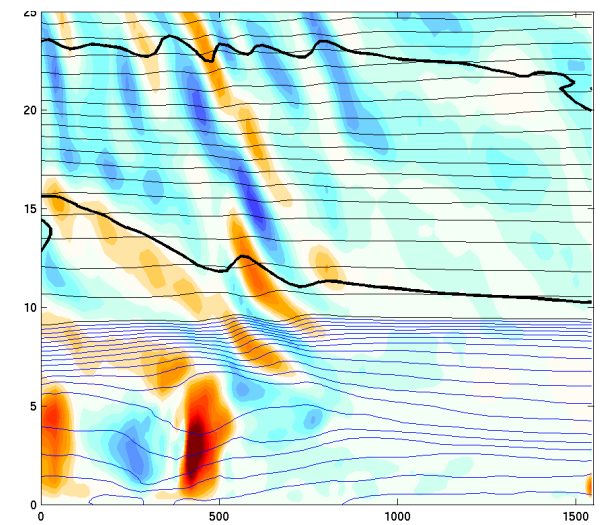
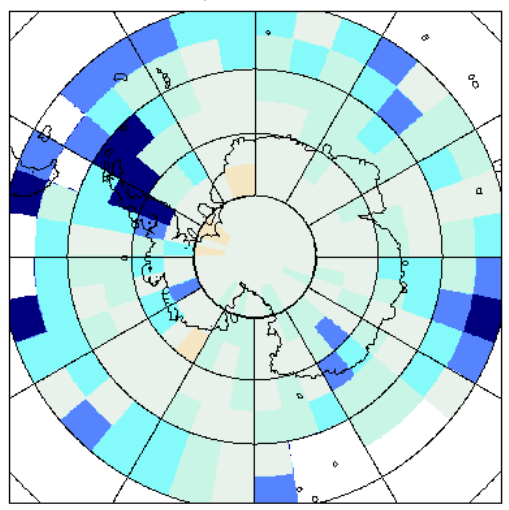
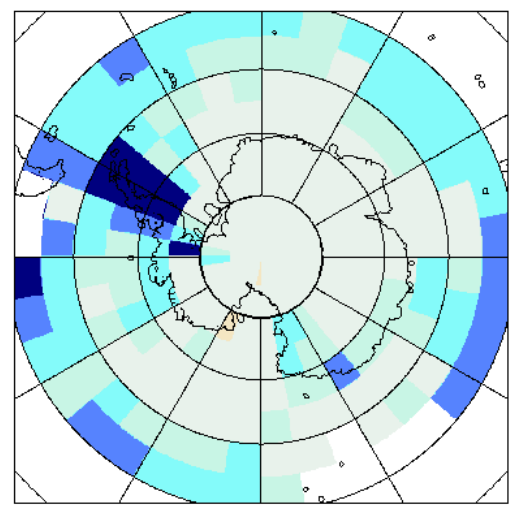


# Gravity waves generated from jets and fronts



Riwal Plougonven,  
 Laboratoire de Météorologie Dynamique,  
 Ecole Polytechnique, Palaiseau, France



# Outline

**Early observations**

**Theoretical understanding of jet exit region waves**

**Numerical simulations // balloon observations**

**Case study**

**Conclusion and discussion**

(Questions, issues in green)

## Favorable configuration for IGW near jets and fronts

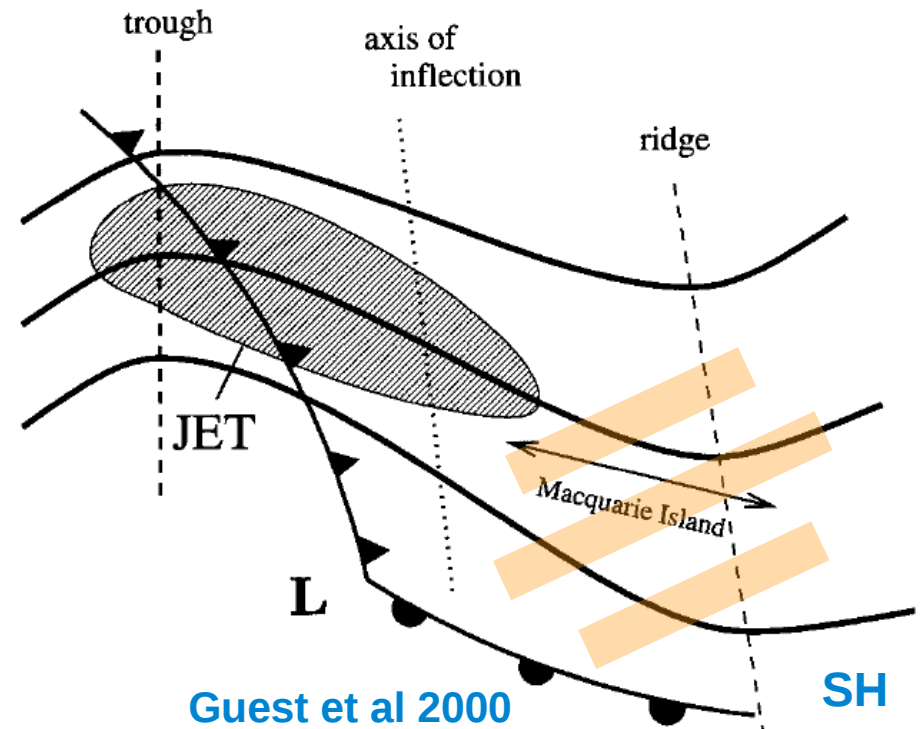
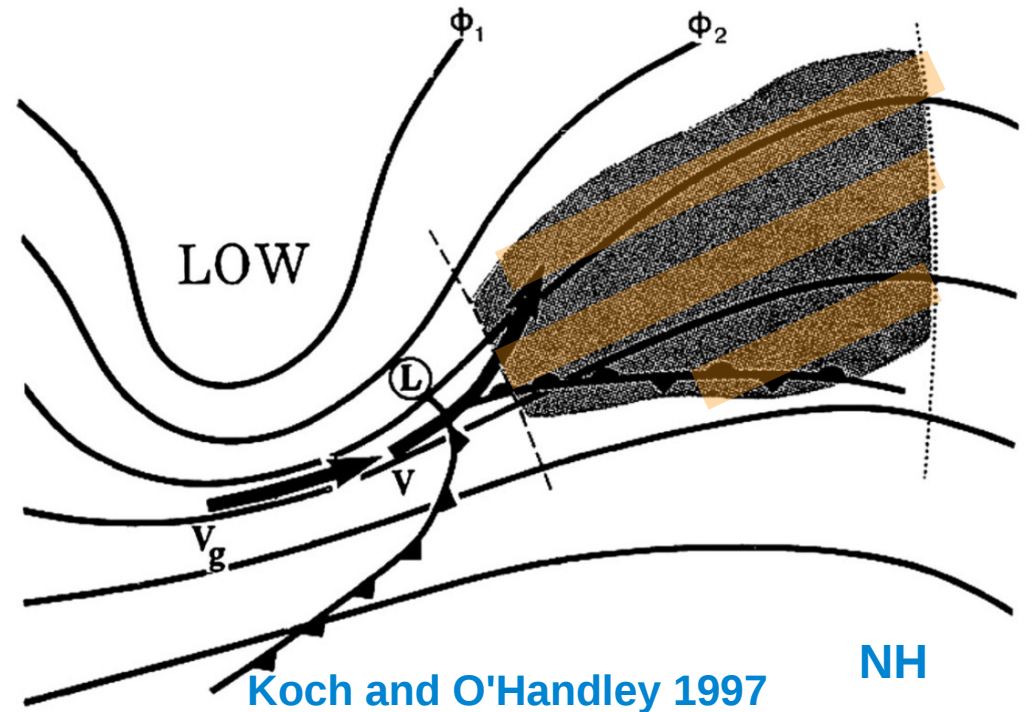
Identified from observational case studies (*Uccellini & Koch 87...*)



Region with IGW

### Questions, issues :

- 1- generation / propagation ?
- 2- only configuration ?
- 3- how important relative to other sources ?
- 4- quantities to describe ?



### 3- Importance of jet/front generated waves relative to other sources ?

Analysis of small-scale motions measured by aircraft

Variations are related to features of the flow

Windspeed and Temperature Variance

JFK-FRA 27 FEBRUARY 1979 2353-0552 UT

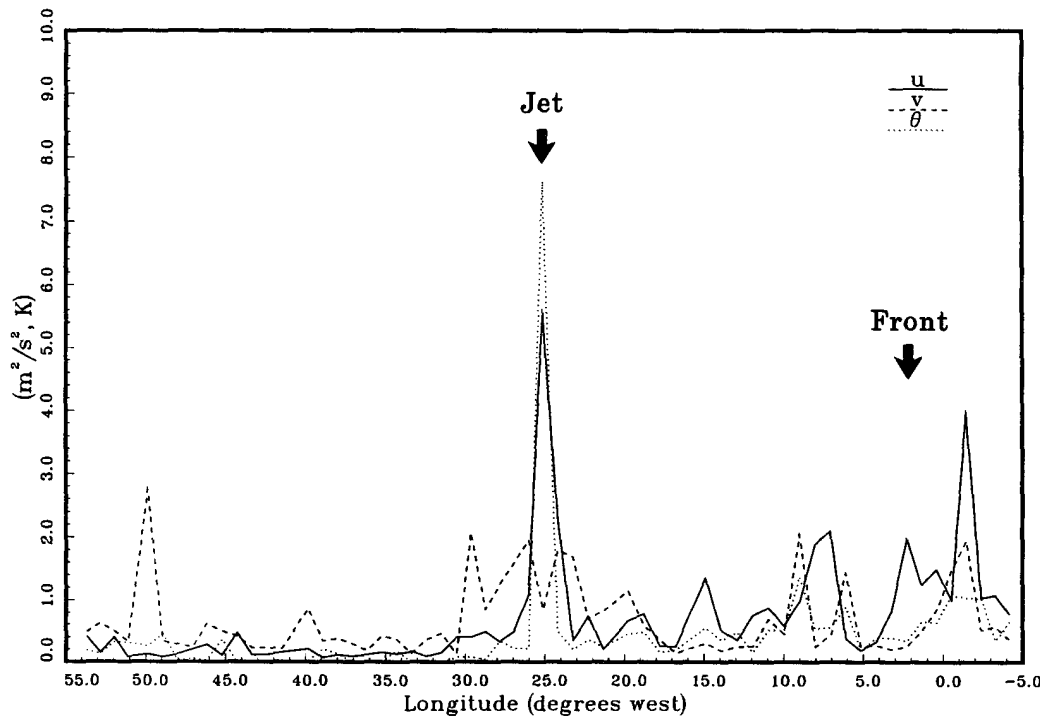
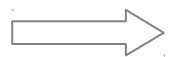
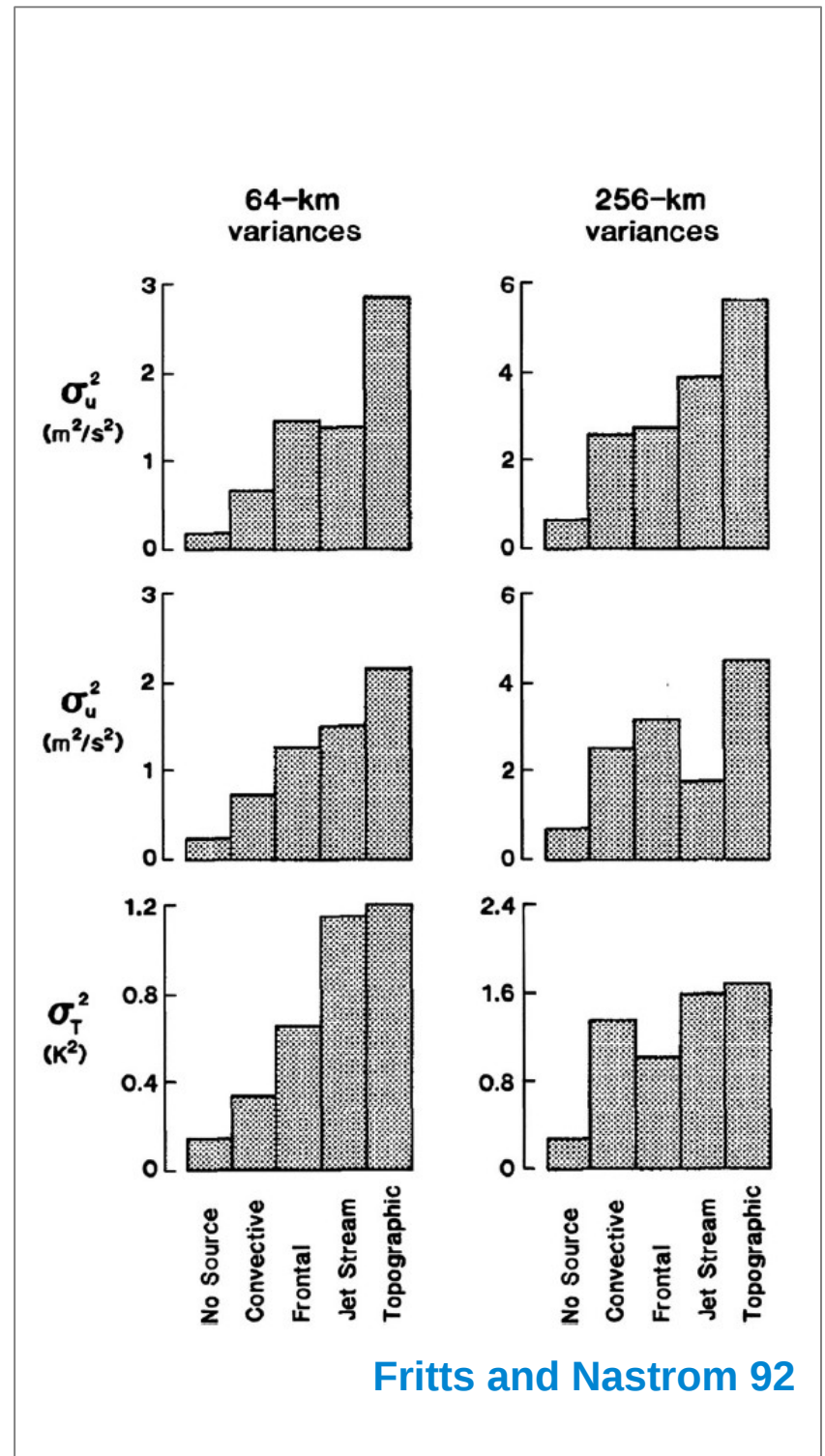


FIG. 12. As in Fig. 11, but for the 64-km variances of  $u$ ,  $v$ , and  $T$  from  $\sim 54^\circ\text{W}$  to  $5^\circ\text{E}$ .



**Jet/front waves of comparable importance to topography**





# Baroclinic instability

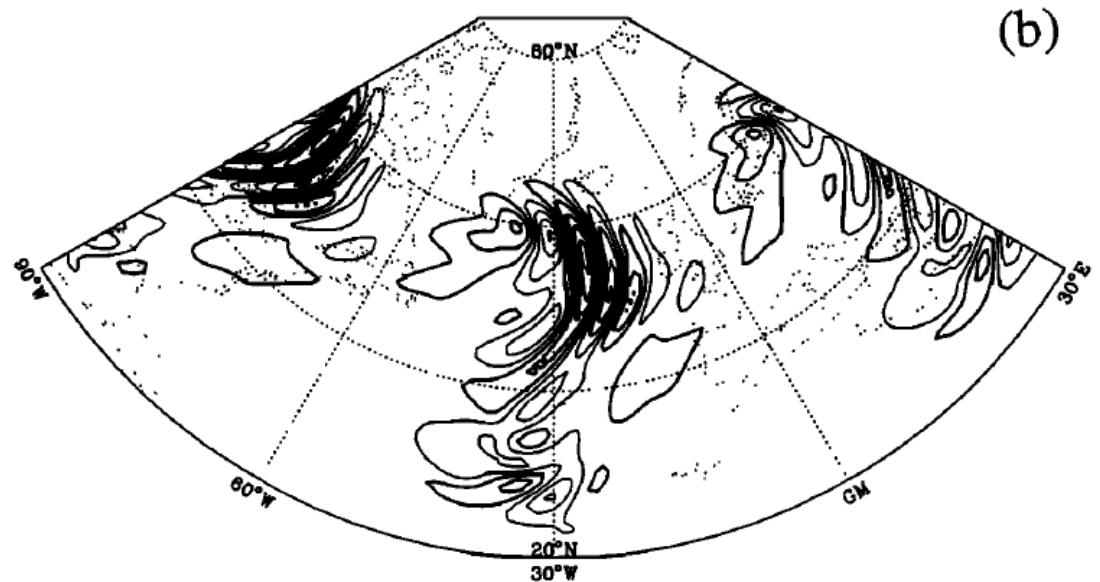
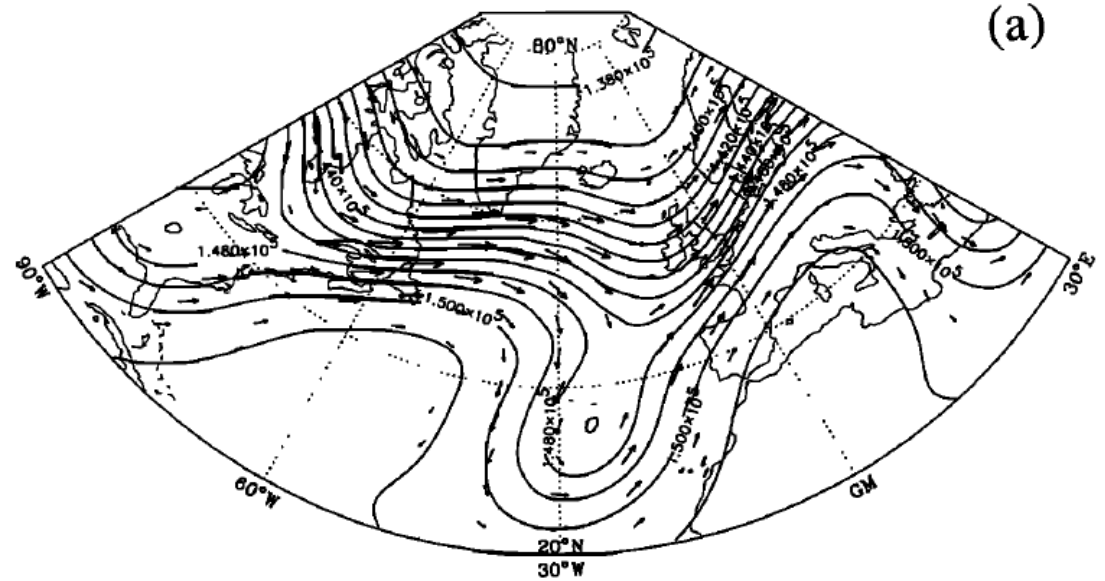
Numerical simulations of an idealized baroclinic instability

GCM with no bottom topography

IGW generated by upper-level jet, **downstream of a jet streak**

Advection of the waves into the rest of the jet (IGW...)

**Generation argued to be well simulated**, yet simulations sensitive to resolution



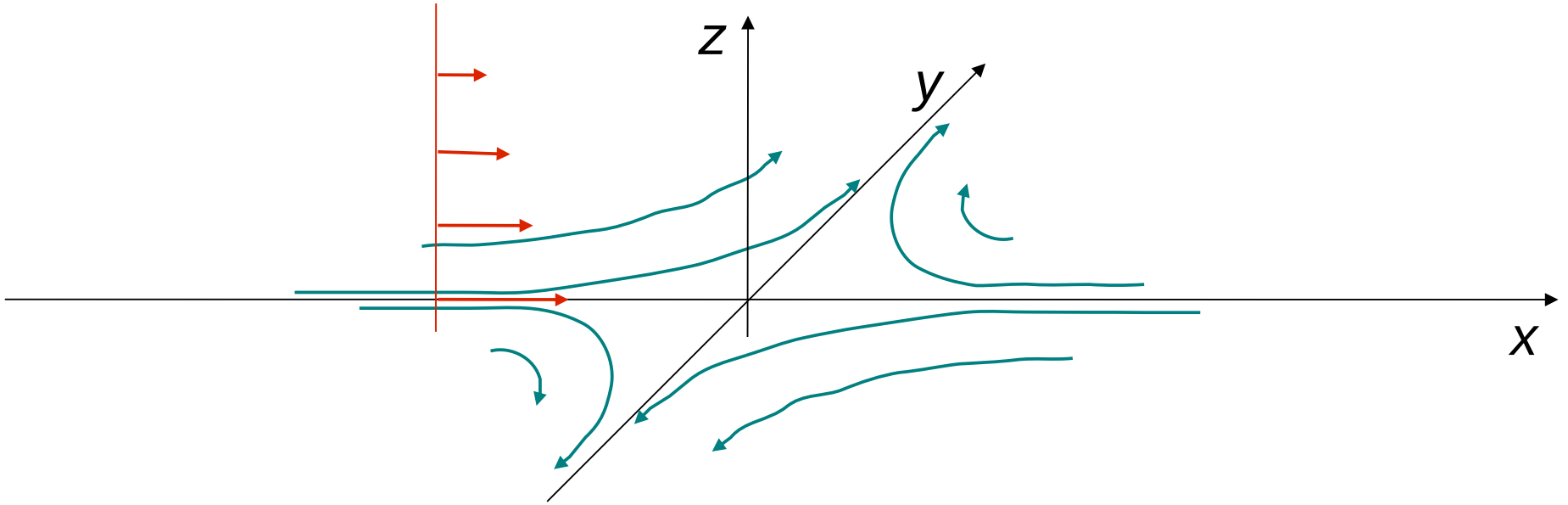
O'Sullivan & Dunkerton 1995

(Bush & Peltier 1995)

Zhang 2004, Plougonven & Snyder 2007

# On the occurrence in jet exit region


Propagation of a gravity wave in a **deformation flow** + **vertical shear** :



$$U(x, z) = -\alpha x + \gamma z,$$
$$V(y, z) = \alpha y + \beta z,$$


$$k = k_0 e^{\alpha t},$$

$$l = l_0 e^{-\alpha t},$$

$$m = m_0 + \frac{\gamma}{\alpha} k_0 (1 - e^{\alpha t}) + \frac{\beta}{\alpha} l_0 (e^{-\alpha t} - 1)$$


**'Wave-capture'** (Buhler & McIntyre 2004, Badulin & Shrira 1993)

IGW packet behaving like passive tracer:

- **phaselines align with extension axis**

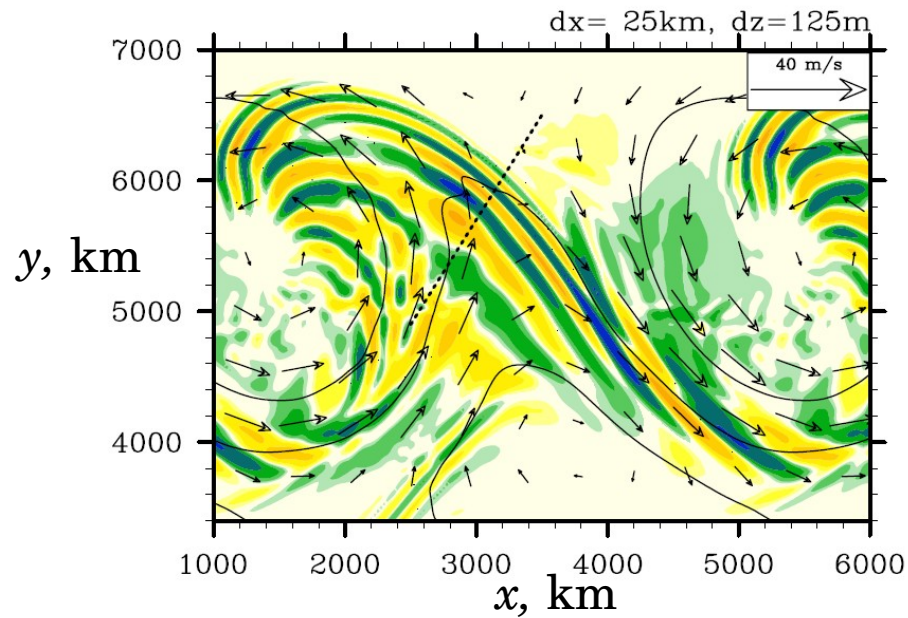
- **wavelengths decrease** exponentially

- 3D orientation ( = **intrinsic frequency**)

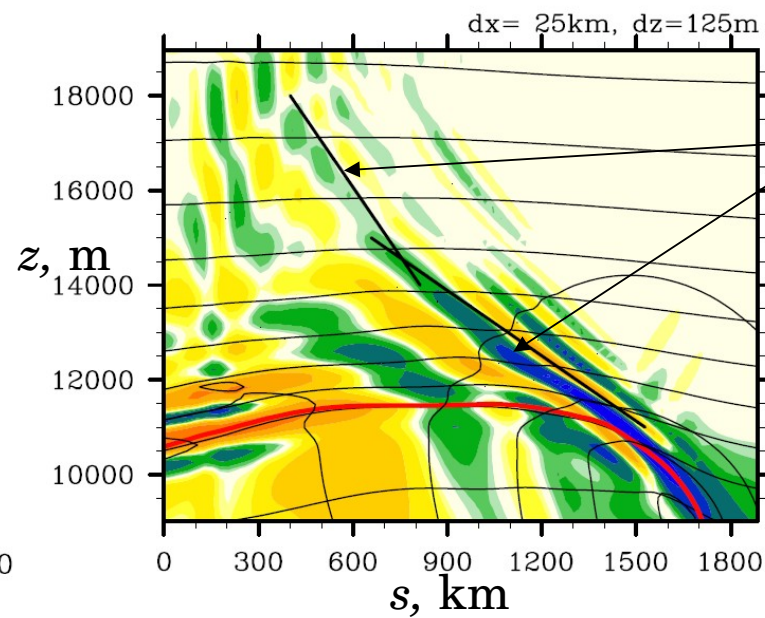
given by large-scale flow:

$$\frac{\lambda_H}{\lambda_z} \rightarrow \frac{|\partial_z U|}{|\partial_x U|}$$

# On the occurrence in jet exit region



Horizontal cross-section of the divergence at  $z=11$  km

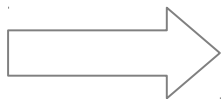
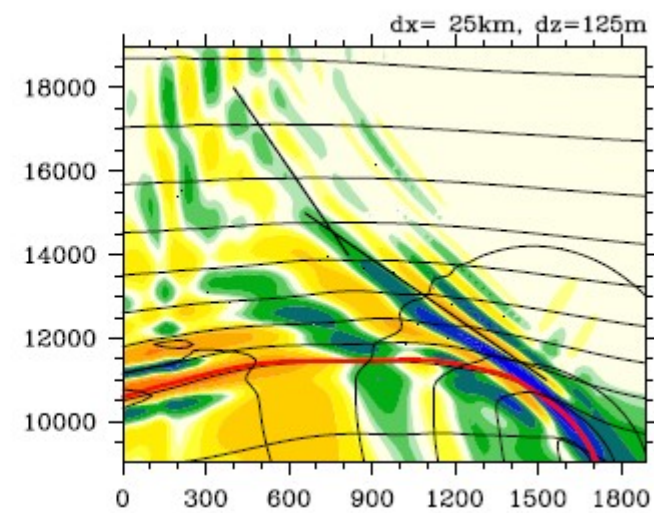
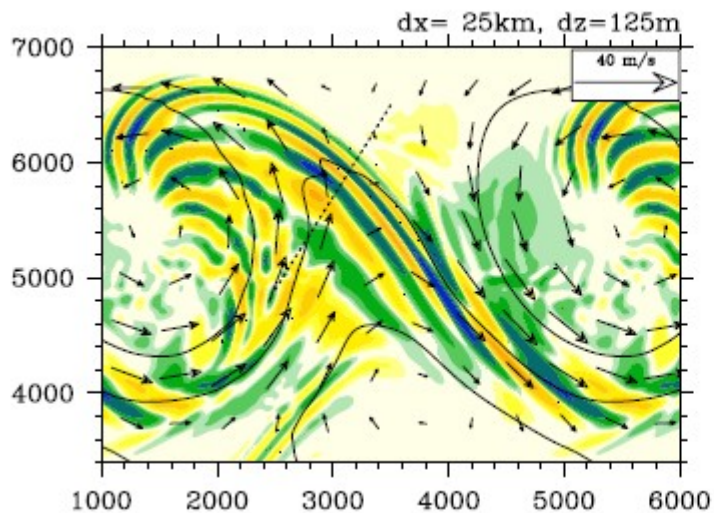
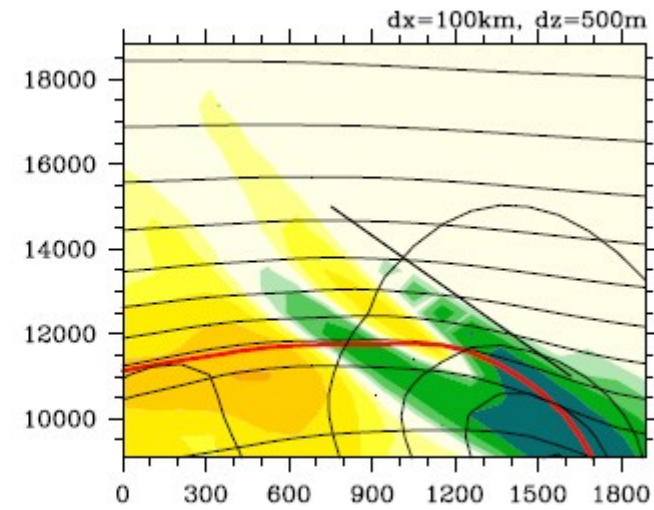
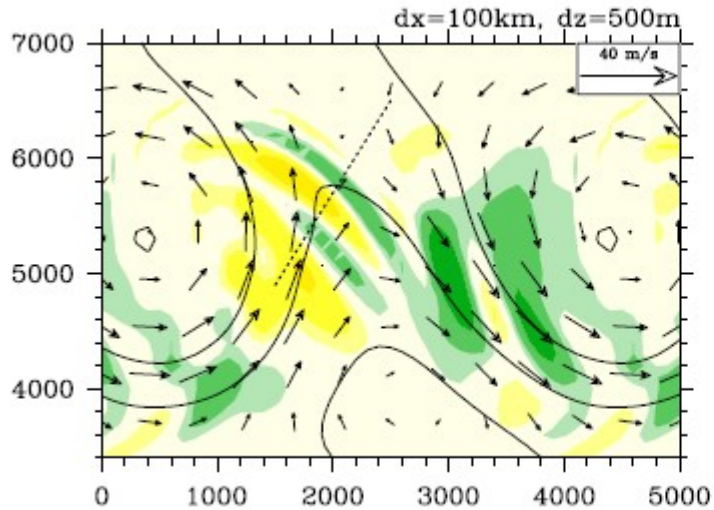


Vertical cross-section of the divergence

Prediction from wave-capture



# On sensitivity to resolution

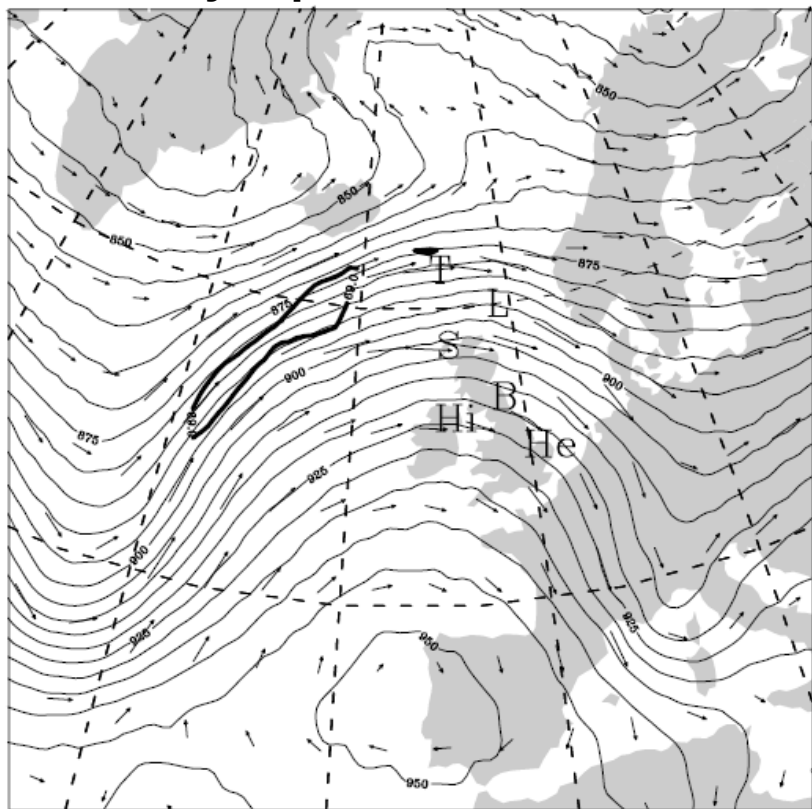


Sensitive to resolution, i.e. Simulations not converged numerically  
But **location, orientation and intrinsic frequency** are insensitive,  
because they are **determined by the large-scale flow**

# On the ability of NWP models to describe GW

Case study comparing IGW described by ECMWF and by 18 radiosondes from FASTEX (1997)

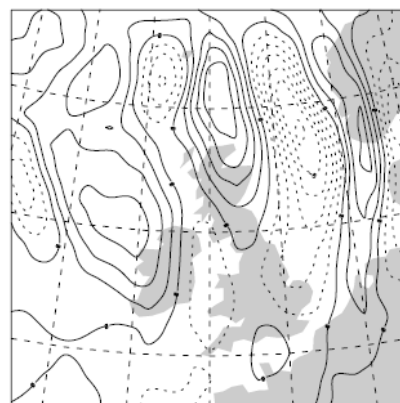
## Synoptic situation



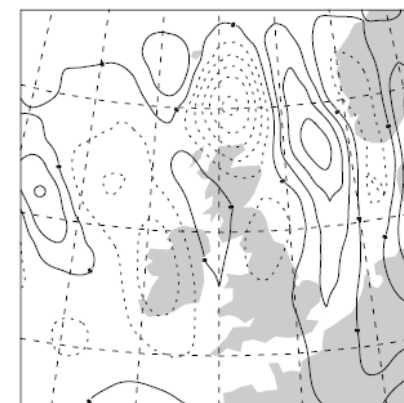
Jet streak ( $-69$  m/s isotach) approaching a ridge of geopotential (300 hPa)

Plougonven & Teitelbaum, 2003

## ECMWF IGW



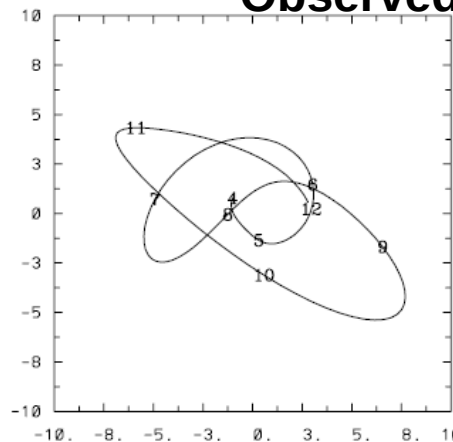
c) Divergence, Feb. 06, 00GMT



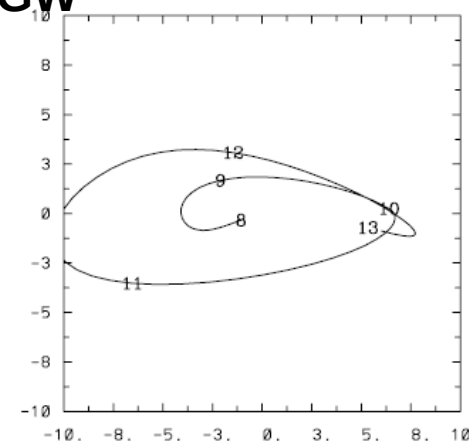
d)  $\theta'$ , Feb. 06, 00GMT

ECMWF described IGW at the right time, right location, right intrinsic frequency and horizontal orientation

## Observed IGW



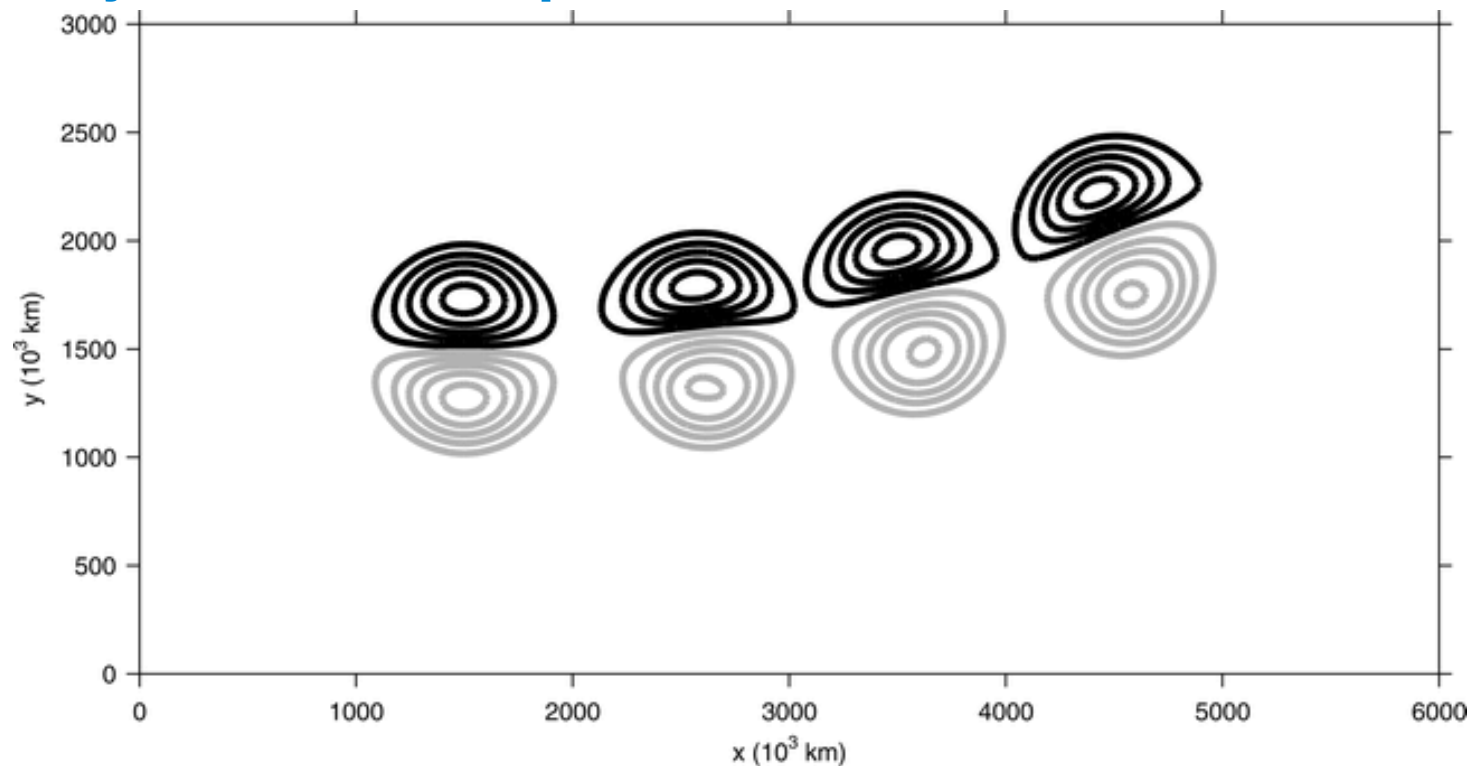
a) Hillsbor., 5/2, 11.15



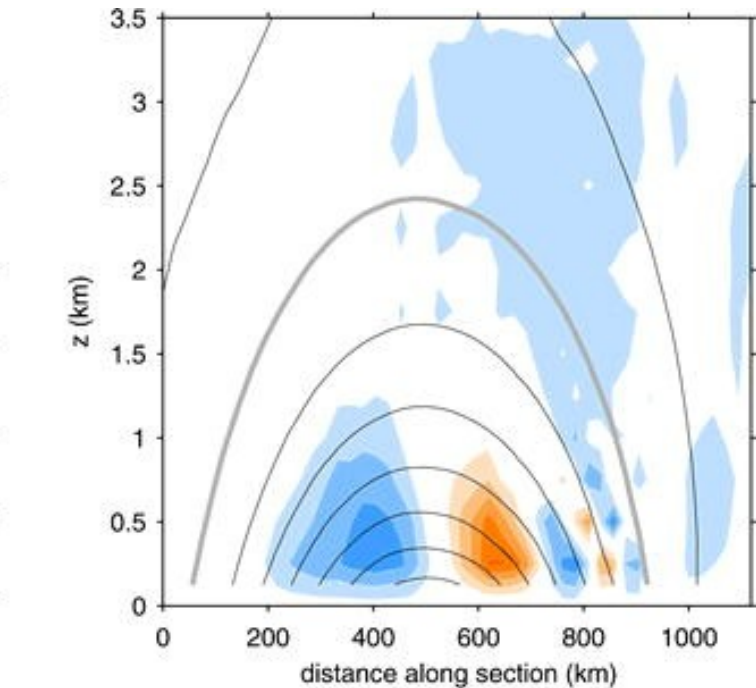
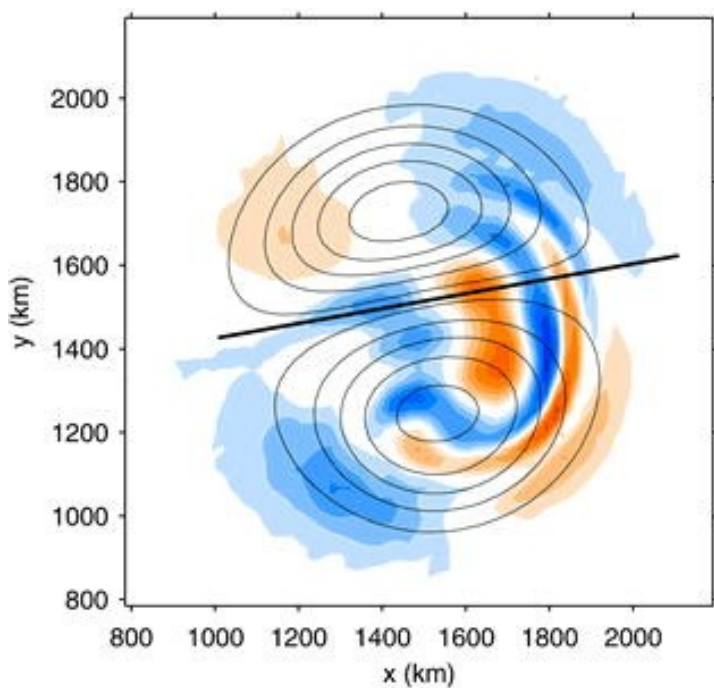
b) Lerwick, 5/2, 17.24

# Simpler model for a jet streak: a dipole

Dipole trajectory  
over 40 days :



Stationary GW  
packet traveling  
with the dipole :



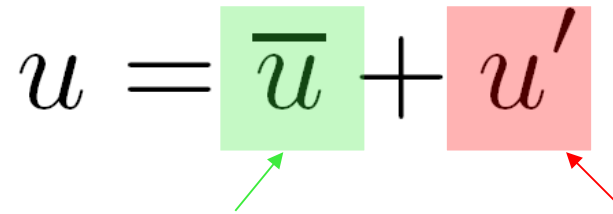
Van Tuyl and Young 1982  
Snyder et al 2007, 2009

# Interpretation

**Mechanism** for the generation:

waves are small perturbations to a dipole that is nearly balanced

1. **Separate** the flow into a balanced part and a perturbation

$$u = \bar{u} + u'$$


Balanced approx.  
of dipole

Small perturbations:  
balanced corrections + IGW

2. **Linearize** about the balanced part

3. Follow **dynamics of perturbations**

homogeneous (instability?)

forced

$$\begin{aligned} \partial_t \bar{u} + \partial_t u' + \bar{\mathbf{u}} \cdot \nabla \bar{u} + \bar{\mathbf{u}} \cdot \nabla u' \\ + \mathbf{u}' \cdot \nabla \bar{u} + \mathbf{u}' \cdot \nabla u' - f \bar{v} - f v' + \partial_x \bar{\phi} + \partial_x \phi' = 0 , \end{aligned}$$

$$\partial_t u' + \bar{\mathbf{u}} \cdot \nabla u' + \mathbf{u}' \cdot \nabla \bar{u} - f v' + \partial_x \phi' = \mathcal{F}_u + \mathcal{O}(u'^2) ,$$

$$\text{where } \mathcal{F}_u = - \left( \partial_t \bar{u} + \bar{\mathbf{u}} \cdot \nabla \bar{u} - f \bar{v} + \partial_x \bar{\phi} \right) .$$



$$\partial_t \bar{u} + \partial_t u' + \bar{\mathbf{u}} \cdot \nabla \bar{u} + \bar{\mathbf{u}} \cdot \nabla u' + \mathbf{u}' \cdot \nabla \bar{u} + \mathbf{u}' \cdot \nabla u' - f \bar{v} - f v' + \partial_x \bar{\phi} + \partial_x \phi' = 0 ,$$

$$\partial_t u' + \bar{\mathbf{u}} \cdot \nabla u' + \mathbf{u}' \cdot \nabla \bar{u} - f v' + \partial_x \phi' = \mathcal{F}_u + \mathcal{O}(u'^2) ,$$

where  $\mathcal{F}_u = - (\partial_t \bar{u} + \bar{\mathbf{u}} \cdot \nabla \bar{u} - f \bar{v} + \partial_x \bar{\phi})$  .

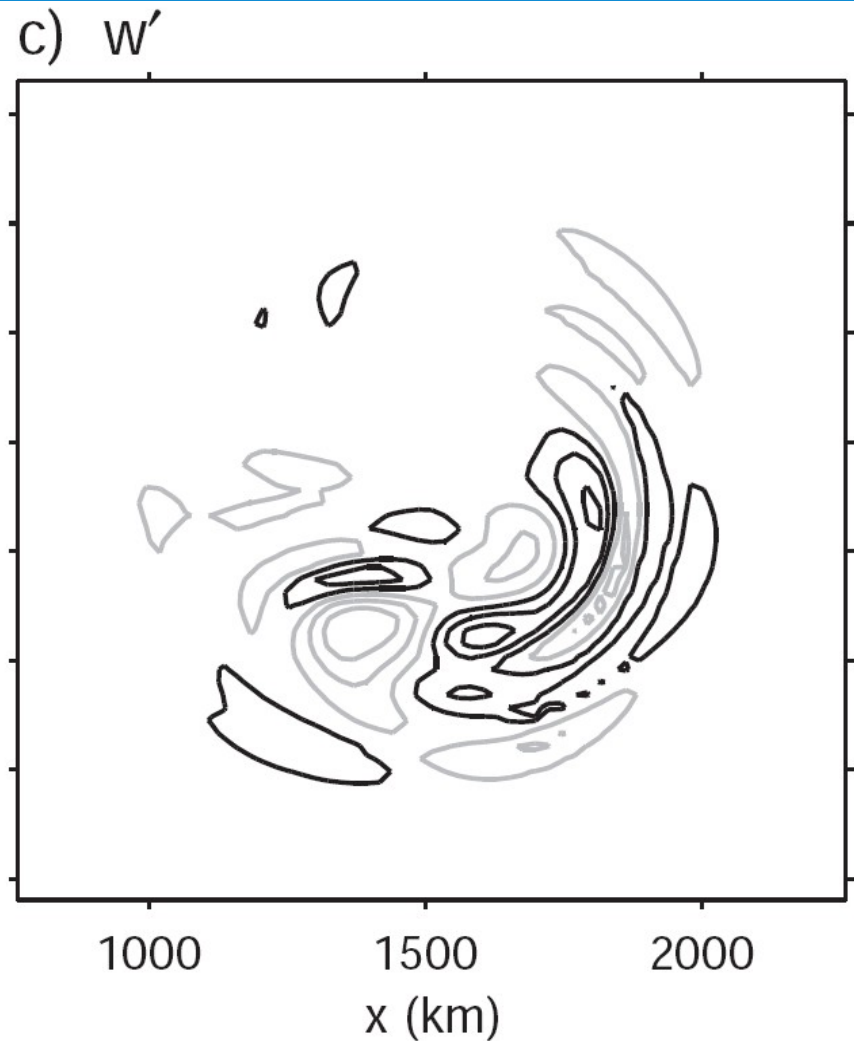
Homogeneous solutions :

$$\partial_t u' + \bar{\mathbf{u}} \cdot \nabla u' + \mathbf{u}' \cdot \nabla \bar{u} - f v' + \partial_x \phi' = 0$$

No growth fast enough to explain waves as the result of an instability,

BUT:

**structures** that emerge in  $w$  very similar to IGW:

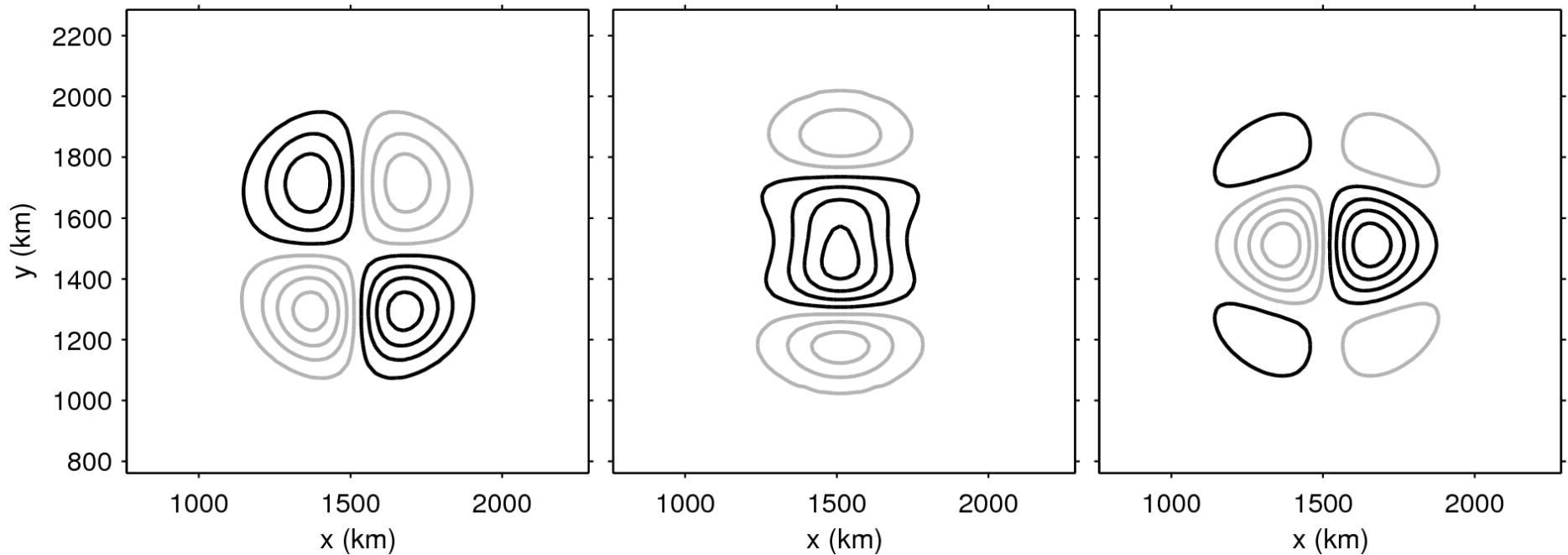


Snyder et al 2009

Forced solutions :

$$\partial_t u' + \bar{\mathbf{u}} \cdot \nabla u' + \mathbf{u}' \cdot \nabla \bar{u} - f v' + \partial_x \phi' = \mathcal{F}_u$$

Forcing terms (known from the QG dipole) :

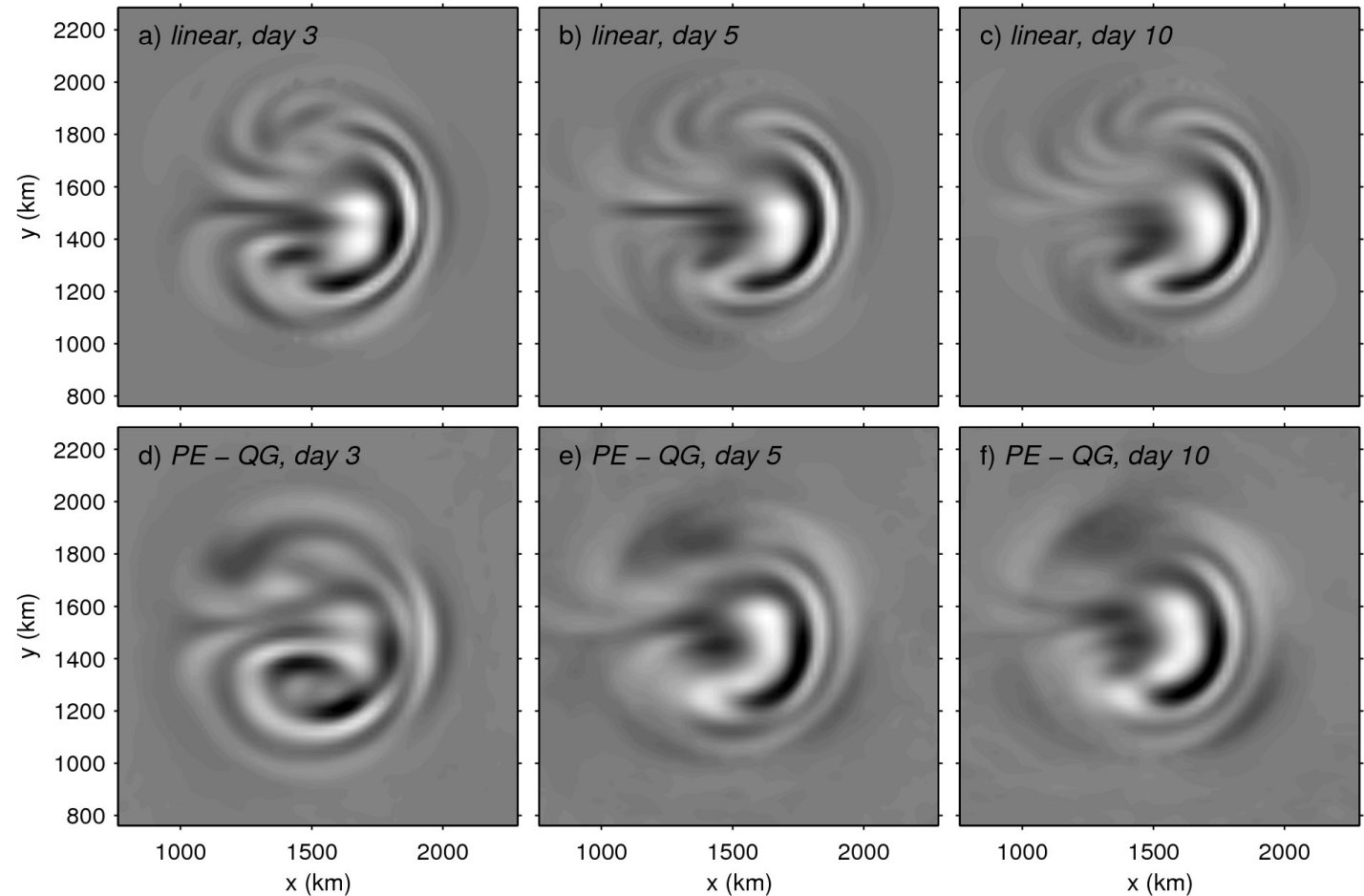


NB all the forcing is large scale, and does not distinguish the front and rear

Forced solutions :

$$\partial_t u' + \bar{\mathbf{u}} \cdot \nabla u' + \mathbf{u}' \cdot \nabla \bar{u} - f v' + \partial_x \phi' = \mathcal{F}_u$$

*W' predicted by  
linear, forced  
calculations:*



*good  
qualitative  
agreement*

# Emission described by renormalization group theory

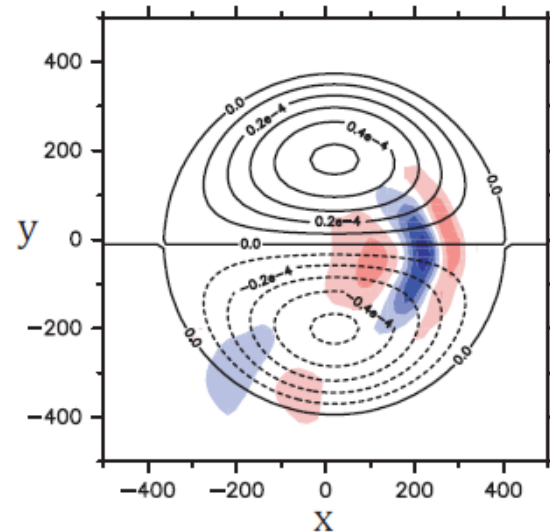
Emission in dipole revisited using **renormalization group theory**

- interpretation of generation mechanism
- interaction of the waves and the dipole

modon, RGE ( $|\Omega| \leq 3$ ,  $|k_1, k_3| \leq 15$ )

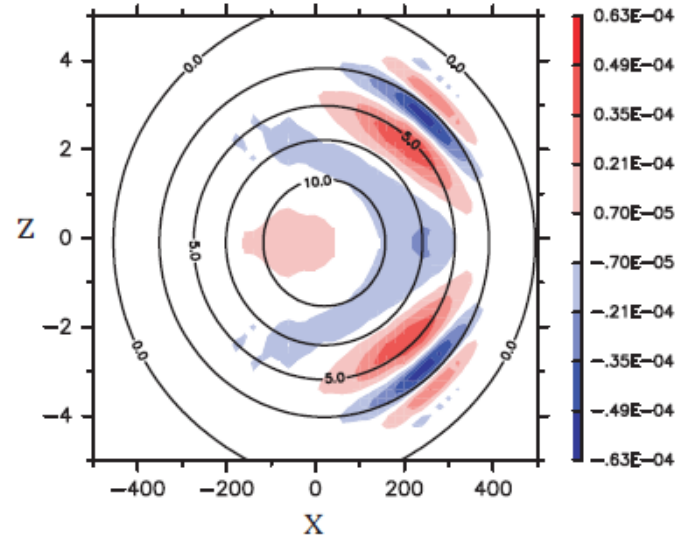
(a)  $\delta^{GW}$ , xy section

$z = 3.0$  km



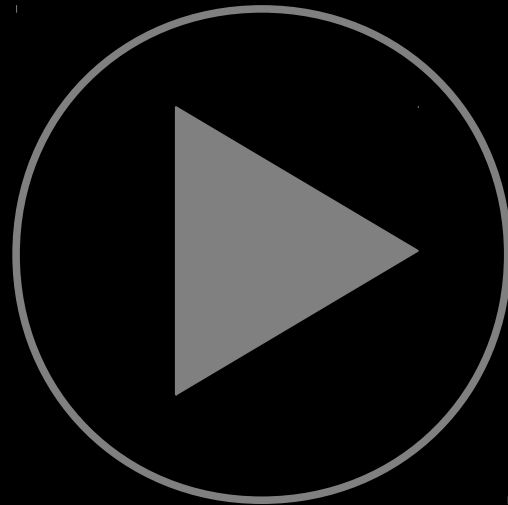
(b)  $\delta^{GW}$ , xz section

$y = 0.0$  km





Toward more realistic flows...  
real case simulations in parallel of balloon observations



# Superpressure balloons

- Constant volume → isopycnic
  - quasi-Lagrangian
  - **intrinsic frequencies**
- Drift in lower stratosphere (~70 – 50 hPa)  
2 months flights on average

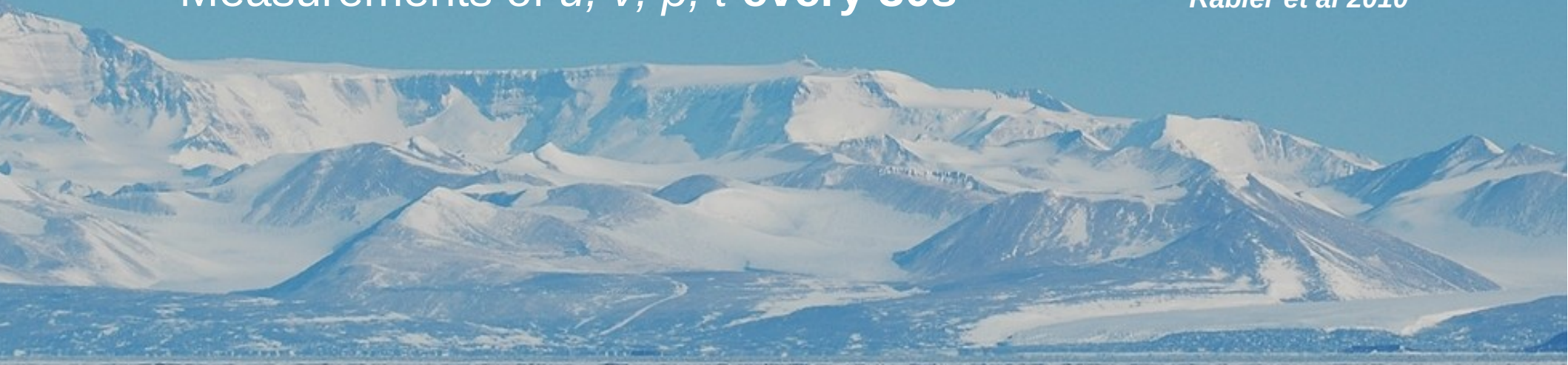


**VORCORE: Sept. 2005 – Feb. 2006, 27 balloons**  
Measurements of  $u$ ,  $v$ ,  $p$ ,  $t$  every 15 min.

*Hertzog et al 2007, 2008*

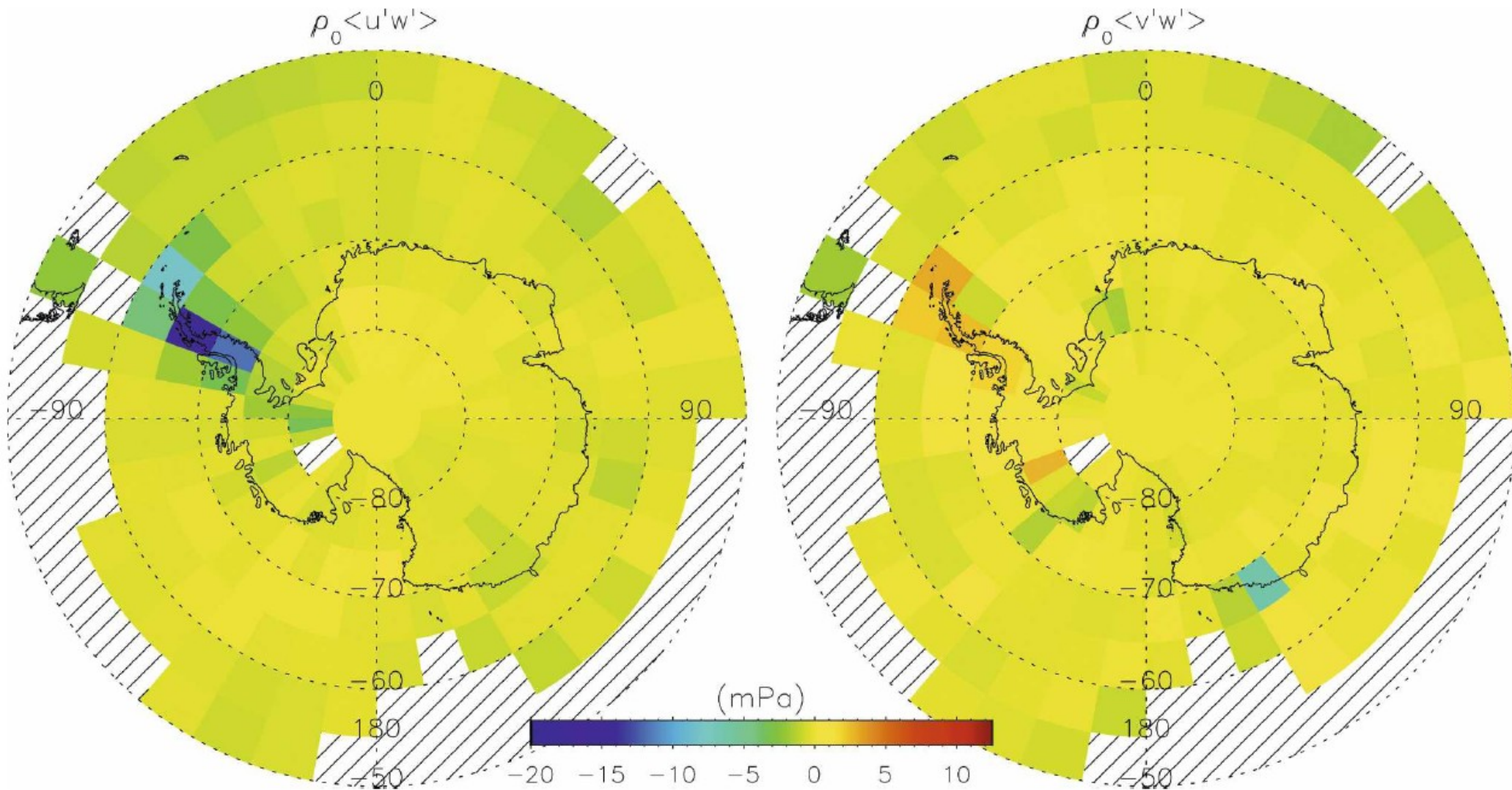
**Concordiasi: Sept. 2010 – Jan. 2011, 19 balloons**  
Measurements of  $u$ ,  $v$ ,  $p$ ,  $t$  every 30s

*Rabier et al 2010*



# Momentum fluxes estimated from VORCORE

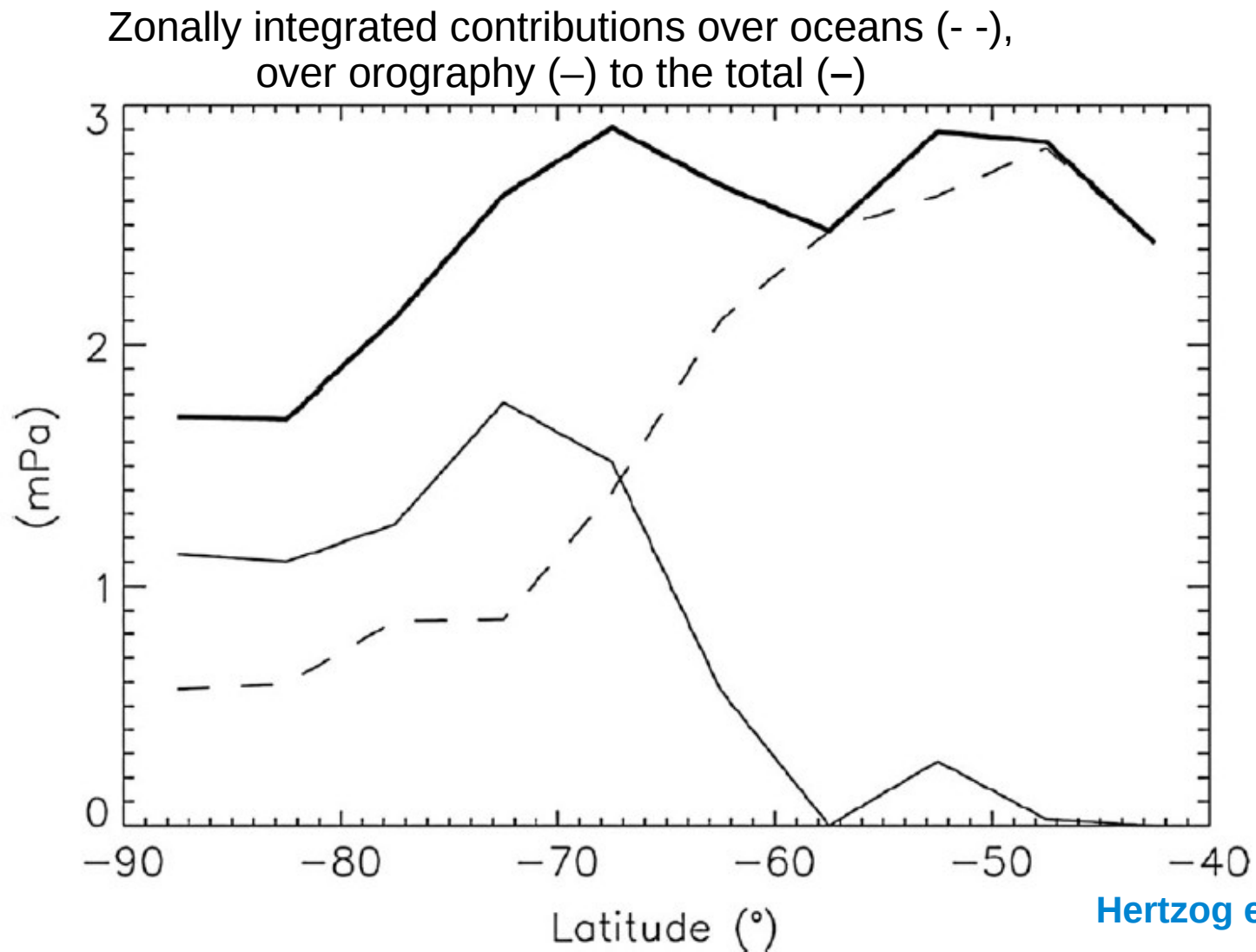
(2005, measurements every 15 min.)



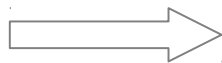
Vincent et al 2007, Hertzog et al 2008

**Orographic hotspot** comes out conspicuously, but...





3



Confirmation of the **importance of non-orographic gravity waves** for momentum fluxes into the middle atmosphere

Role in the late breakup of the Southern stratospheric polar vortex ?

McLandress et al 2012, Jewtoukoff et al 2015

# Mesoscale simulations in parallel of VORCORE

Weather Research and Forecast (WRF) Model (Skamarock et al 2008)

## Domain and resolution :

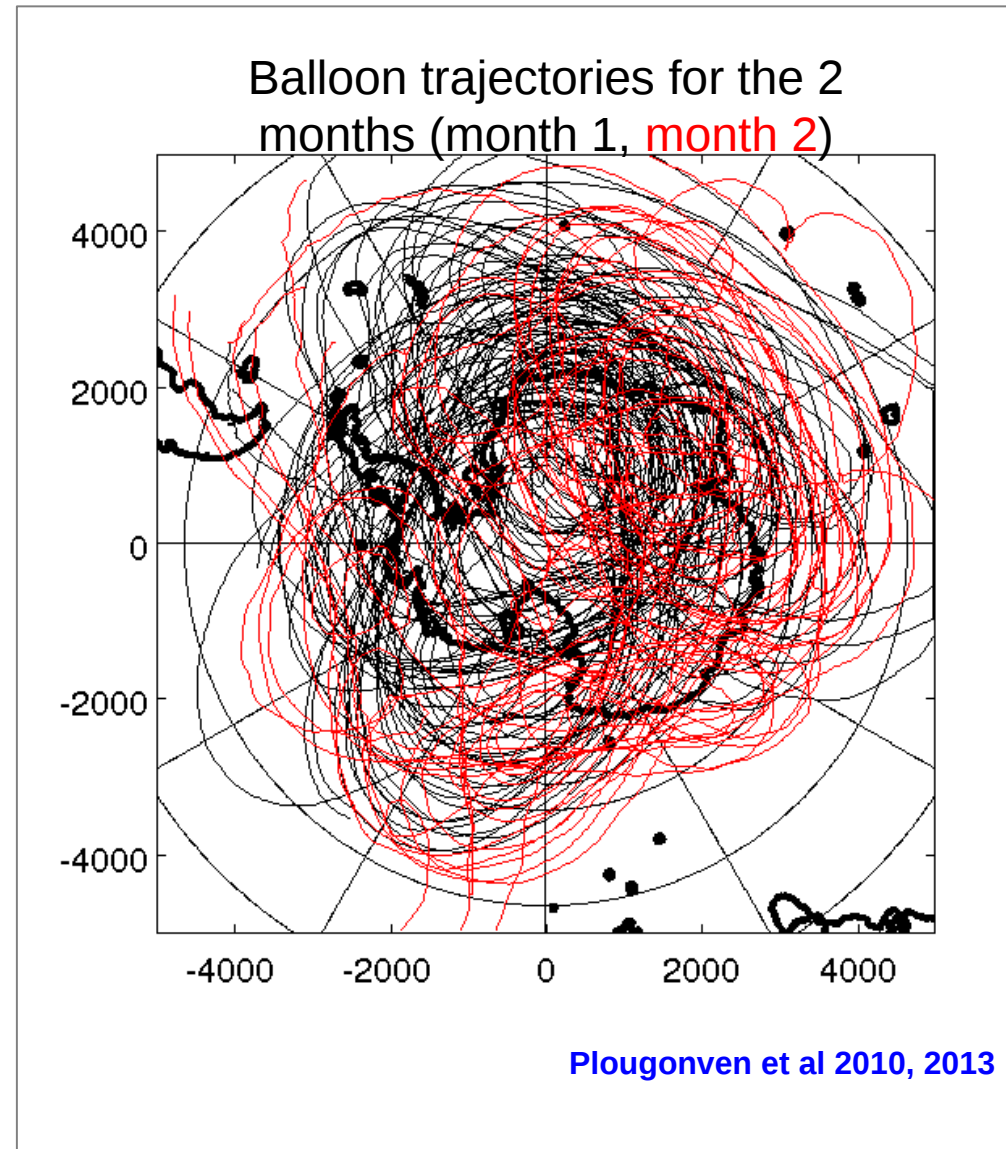
**10 000 km x 10 000 km,  $dx = 20$  km**  
up to 5 hPa with 120 levels

## Time period :

October 21, 2005, 00:00UT

**58 days**

December 18, 2005, 00:00UT

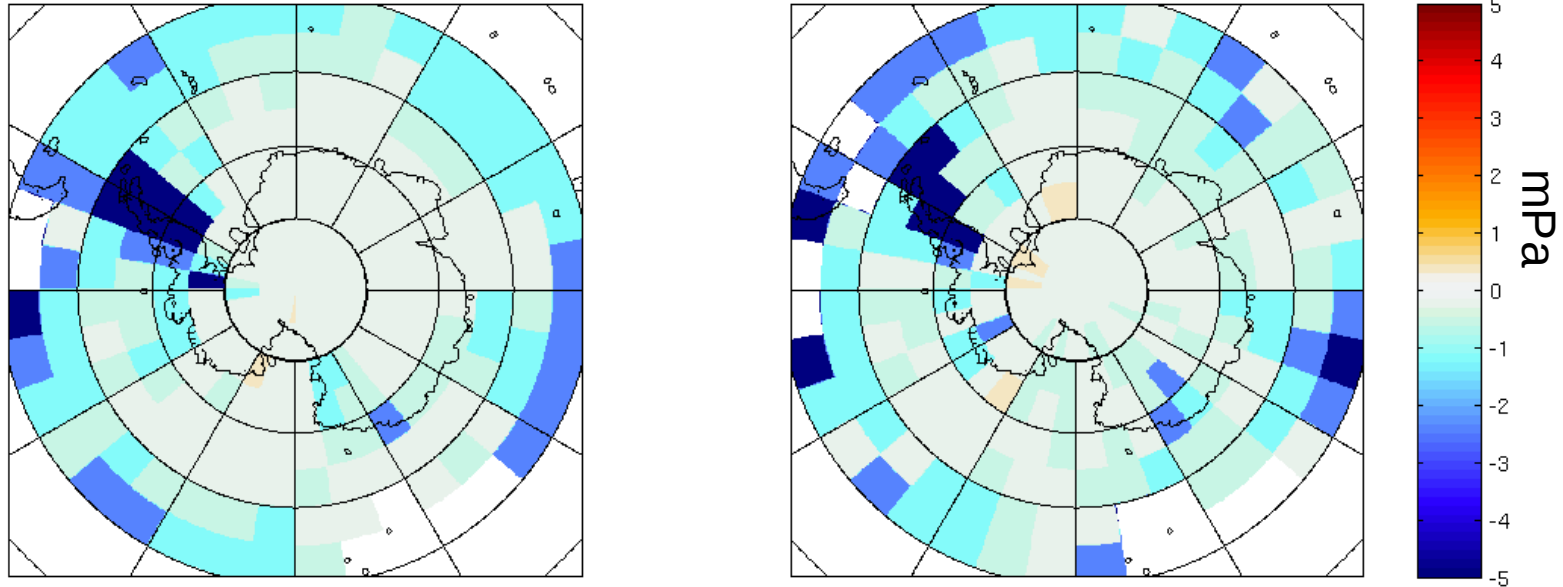




# Mesoscale simulations in parallel of VORCORE

Time-averaged momentum fluxes ( $\langle \rho u' w' \rangle$ ):

Comparison of the mean zonal momentum fluxes in the  
simulations & observations



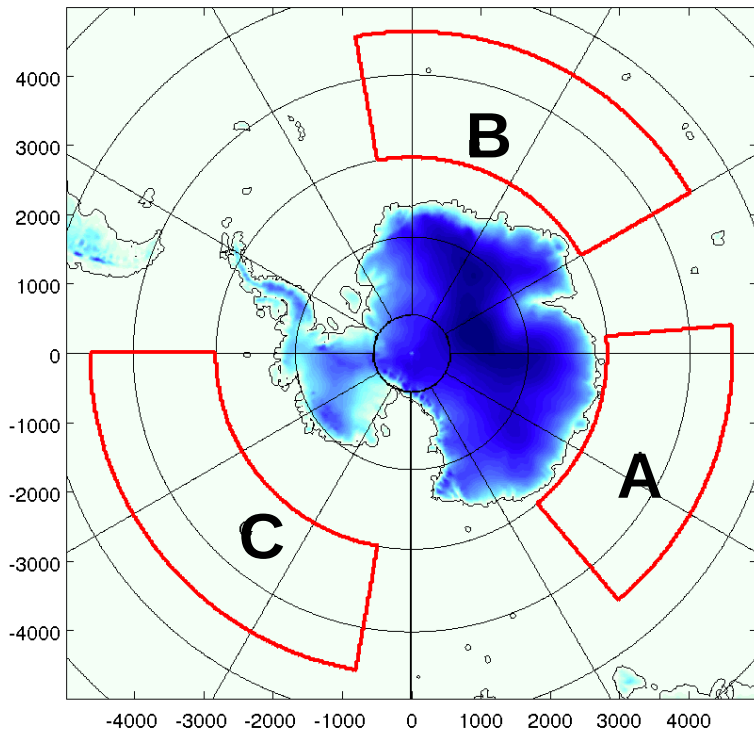
**Good overall agreement,**  
Similar order of magnitude

- maximum over the Antarctic Peninsula
- **comparable structures and amplitude over the ocean**  
(observed average : 0.83 mPa, simulated average : 0.67 mPa)

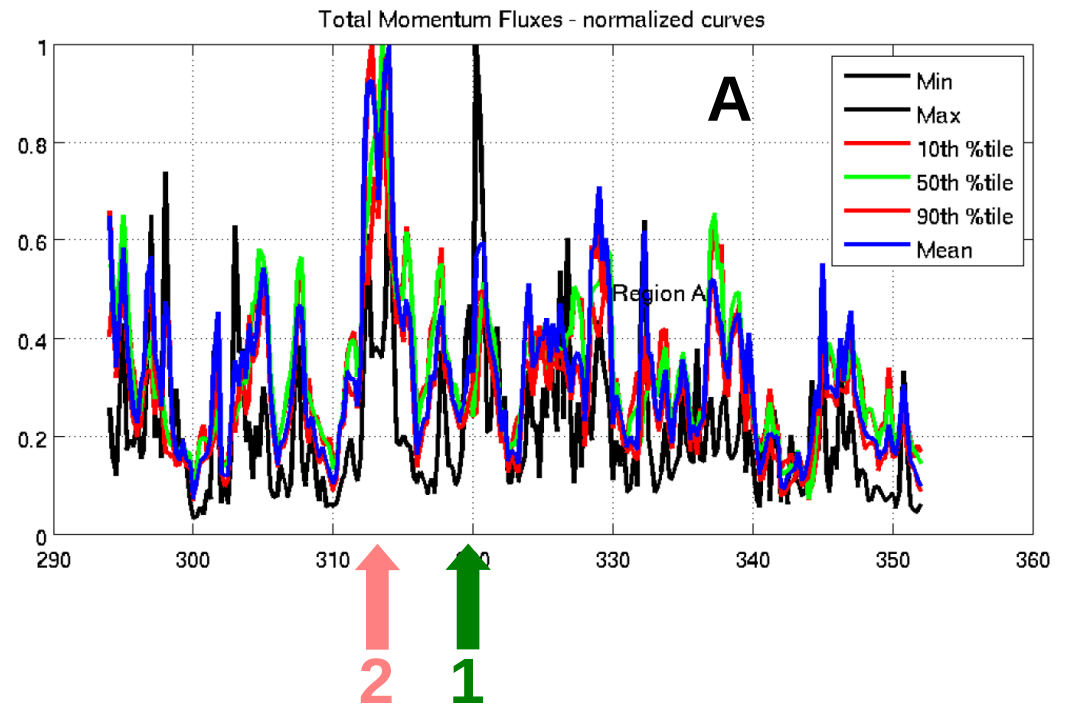
# Case study

## Choice of case studies :

**1.** Identify 3 regions where GW are non-orographic



**2.** Examine time series of mean momentum fluxes over each region

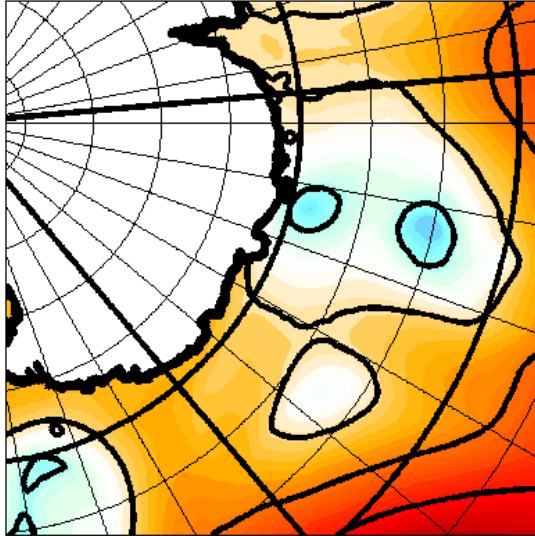


**3.** Choose case studies for intense events :

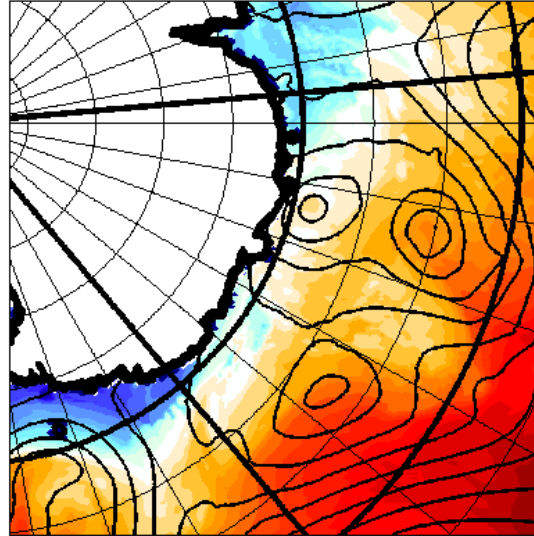
- **Case 1 : day 320**
- **Case 2 : day 313**
- Case 3 : region C, day 296

## Case study

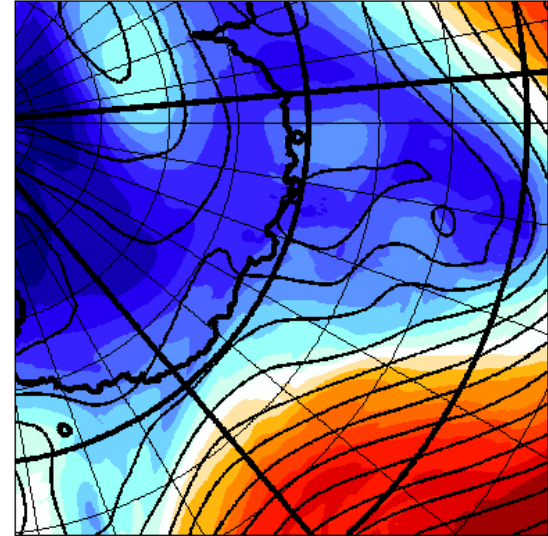
### Meteorological situation – tropospheric flow



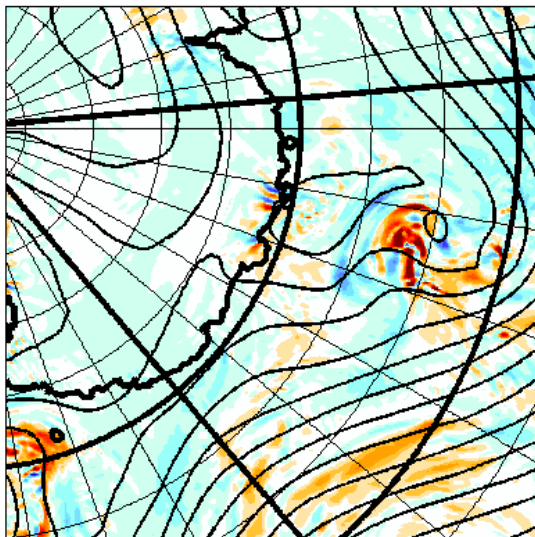
Surface pressure



Surface temperature  
(+ isobars)



Potential temperature  
at  $z=5$  km (+ isobars)



Vertical velocity at  
 $z=5$  km (+ isobars)

November 15, 2005, 18:00

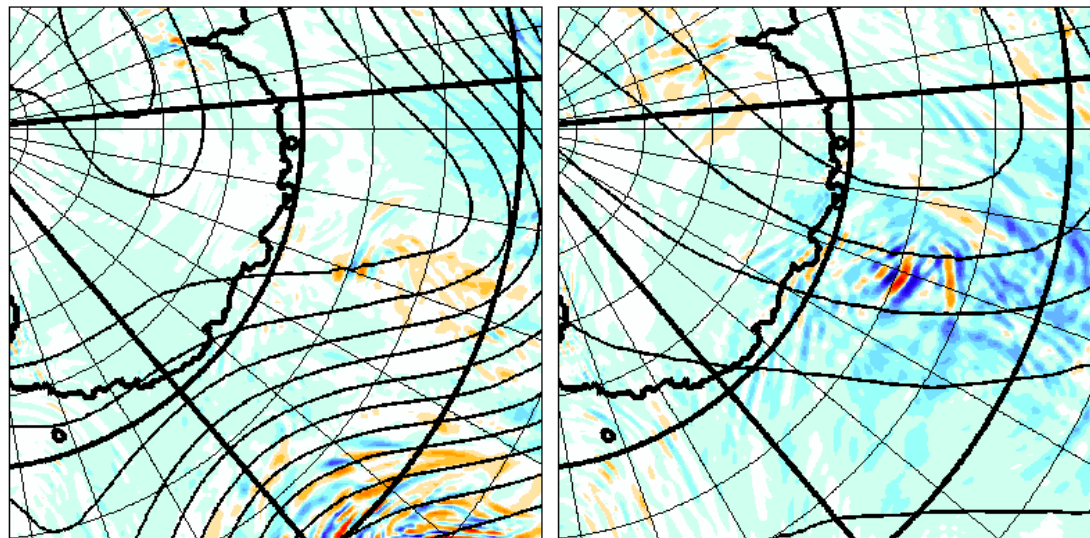
## Case study

# Stratospheric gravity waves

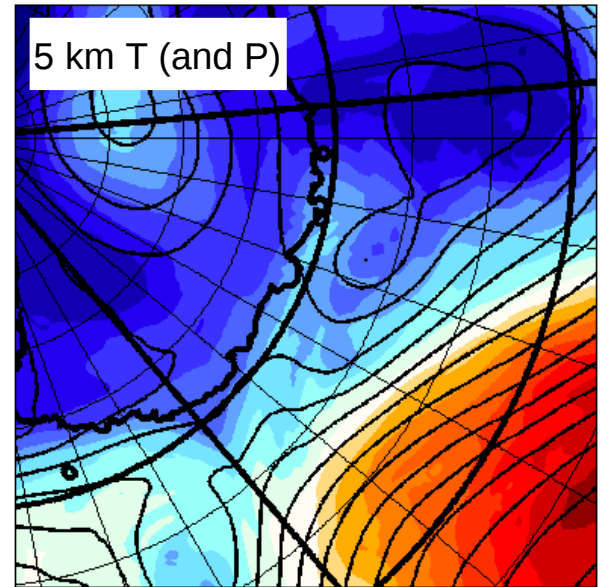
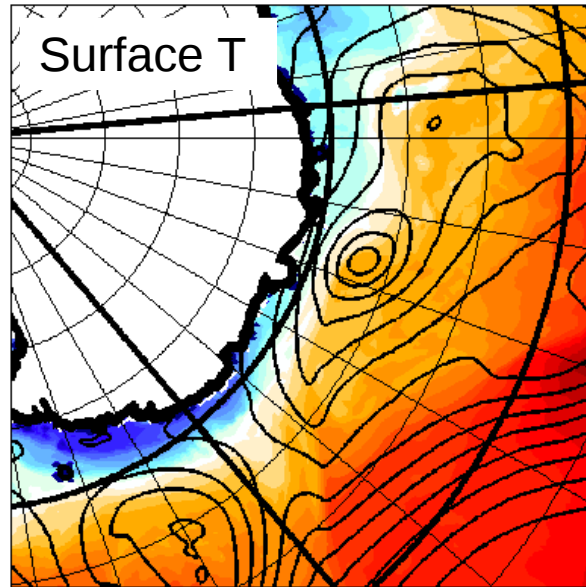
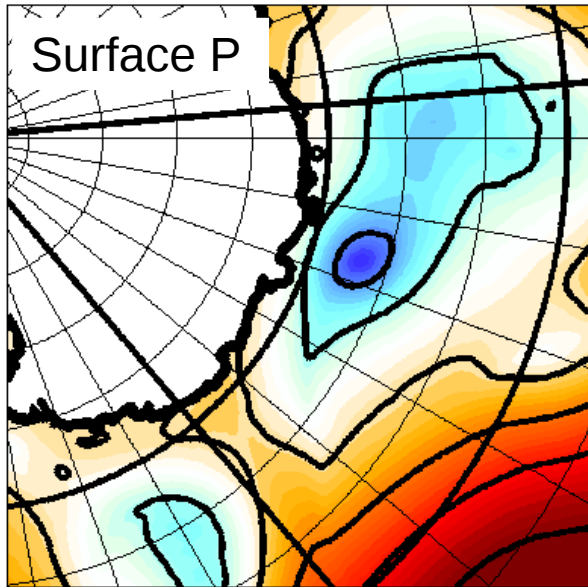
November 15, 2005, 18:00

Vertical velocity at  
 $z=10$  km (+ isobars)

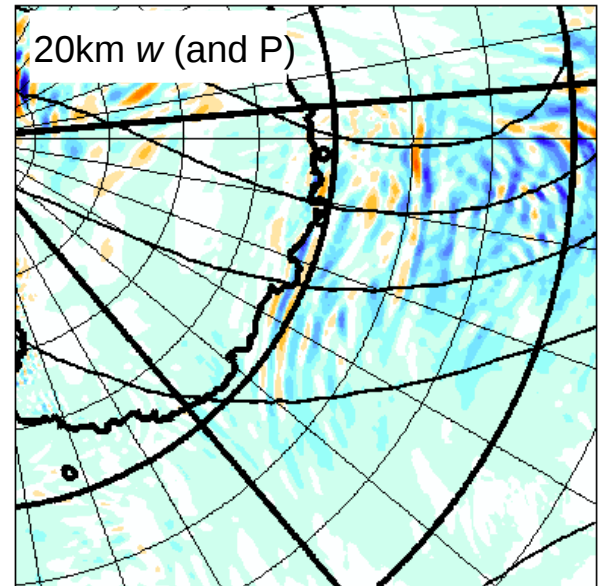
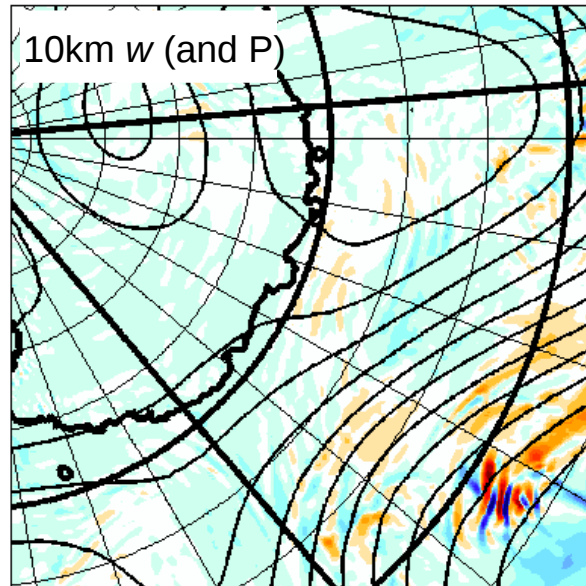
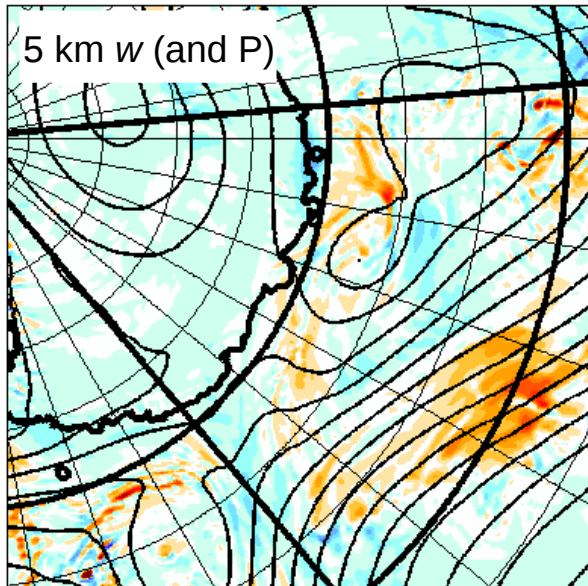
Vertical velocity at  
 $z=20$  km (+ isobars)



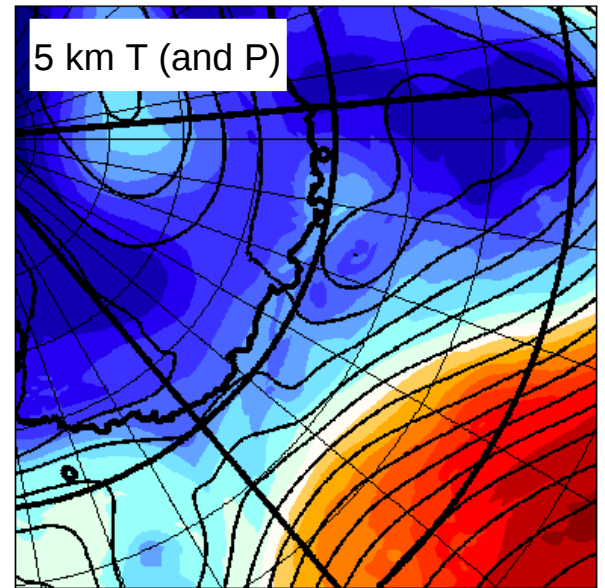
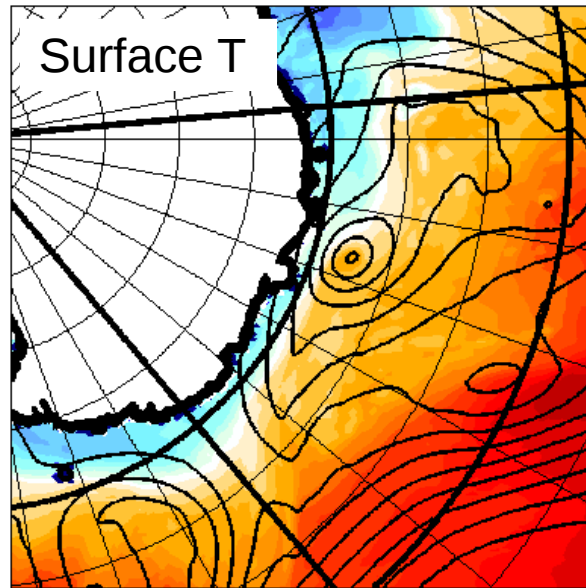
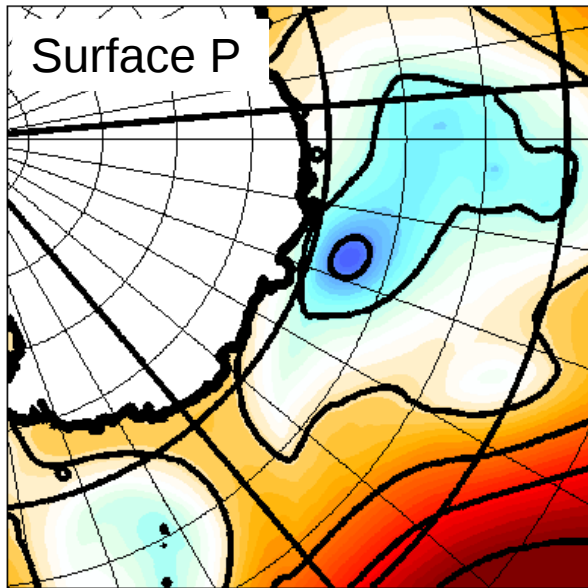




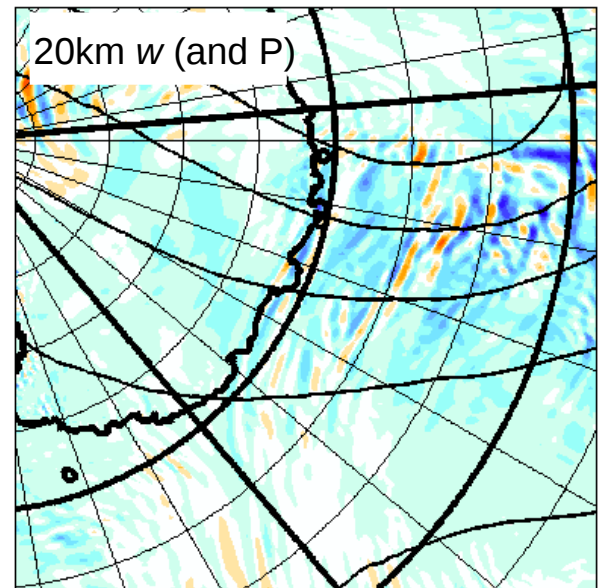
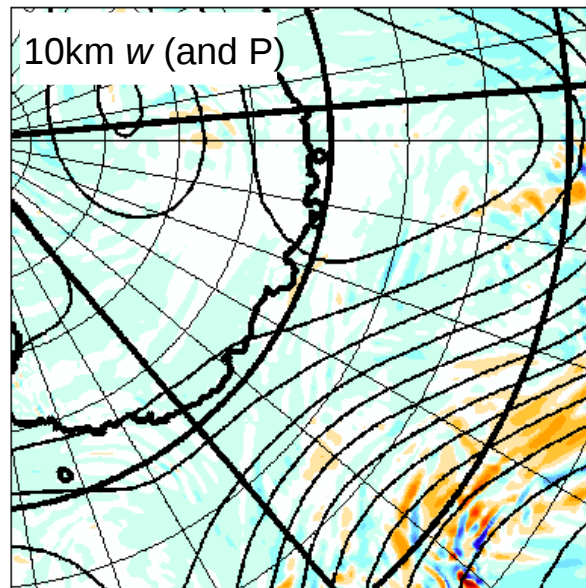
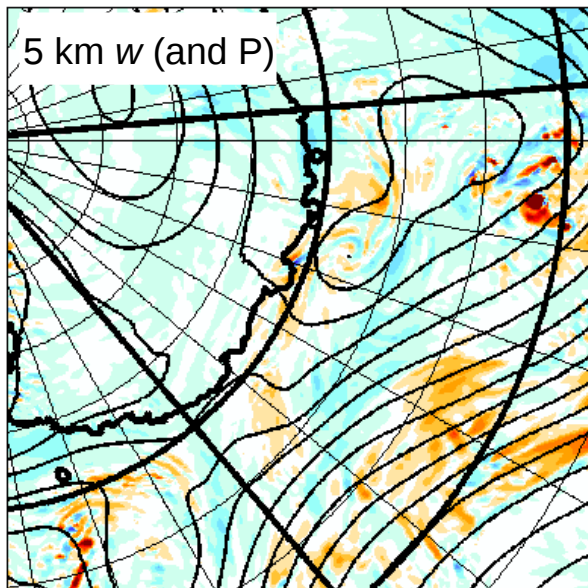
November 14, 2005, 18:00

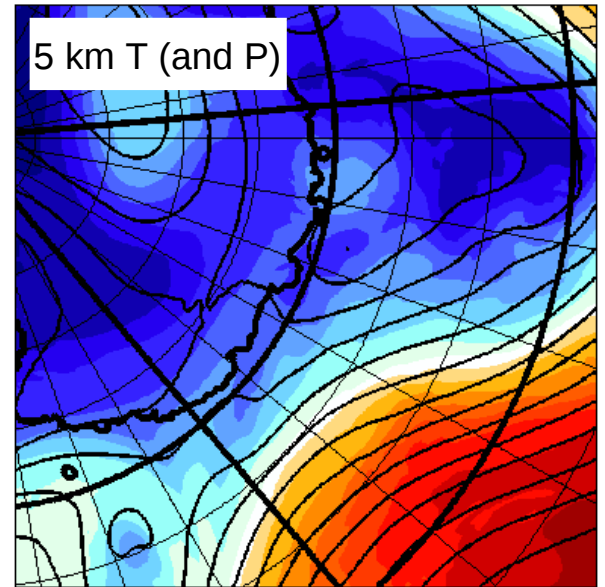
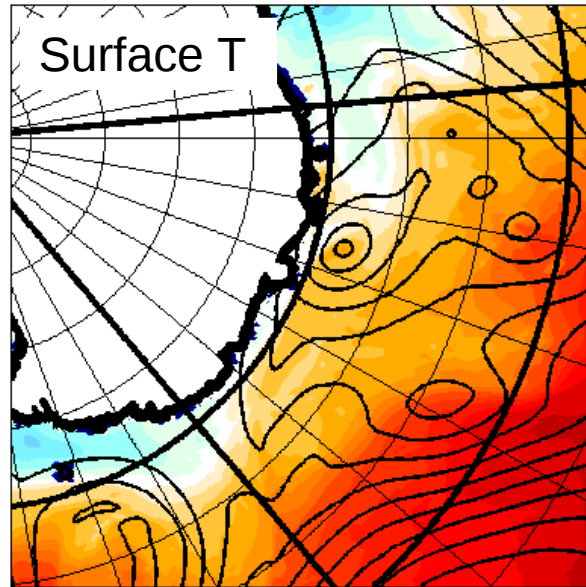
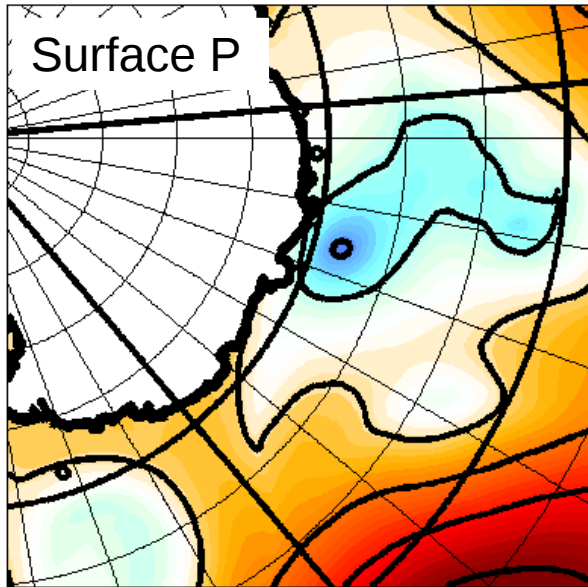




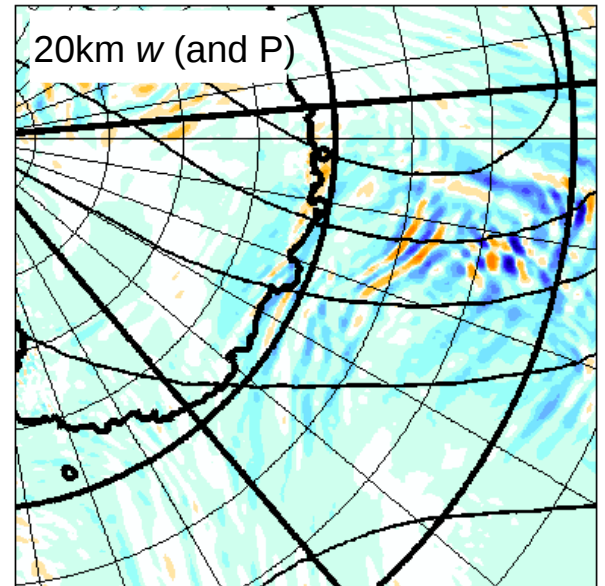
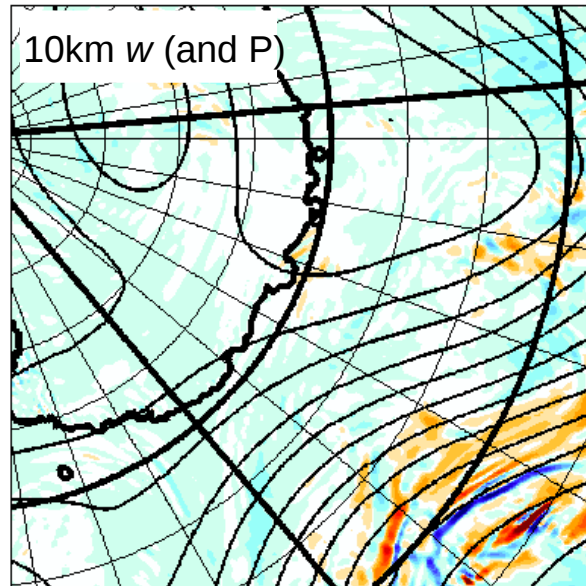
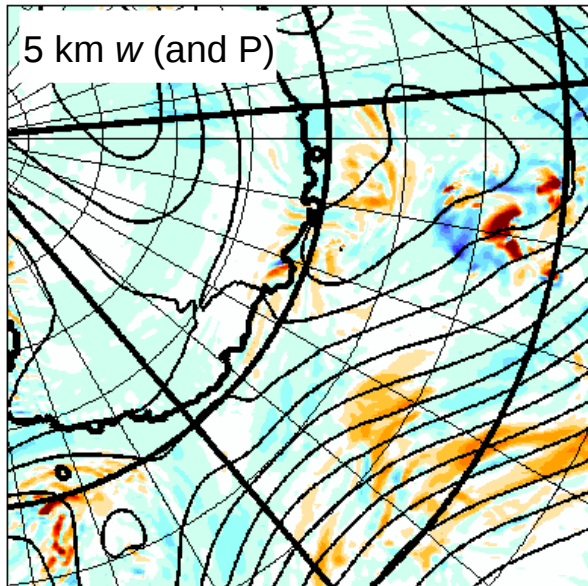


November 15, 2005, 00:00

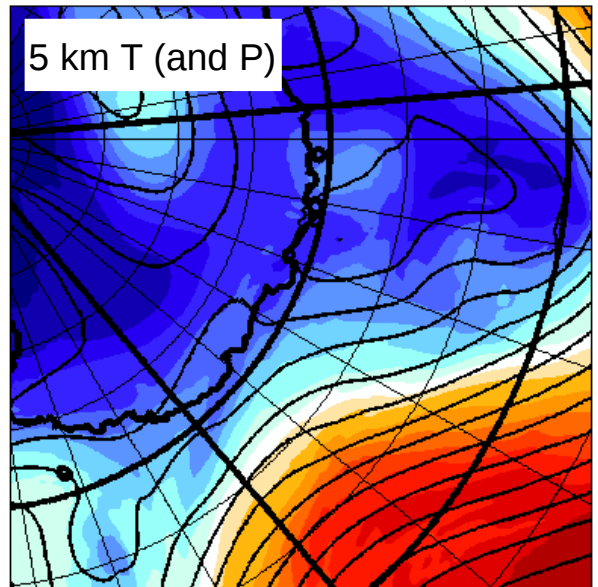
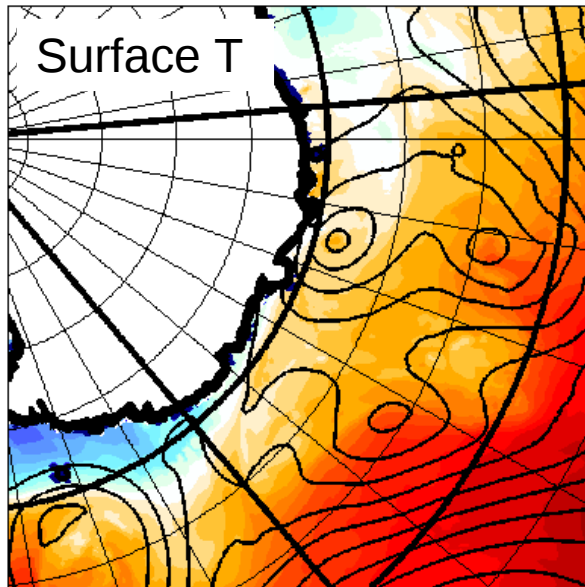
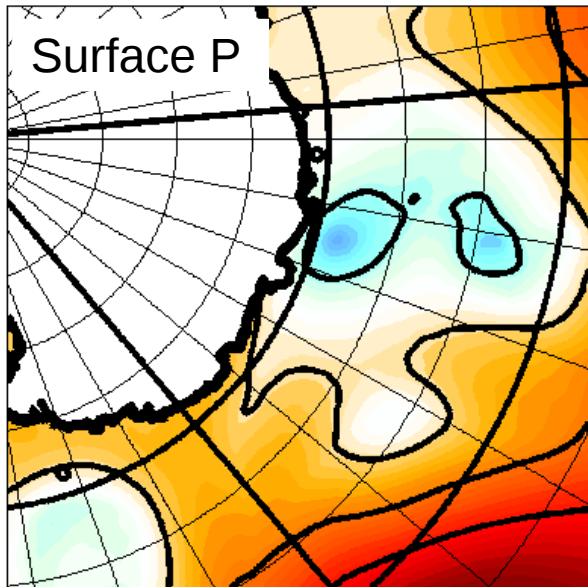




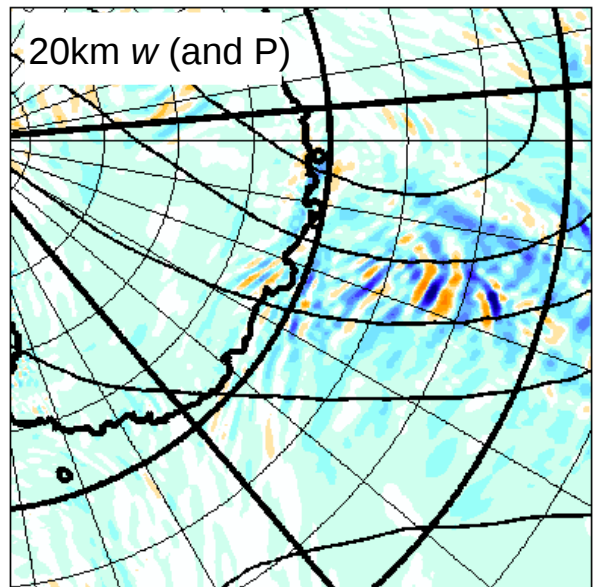
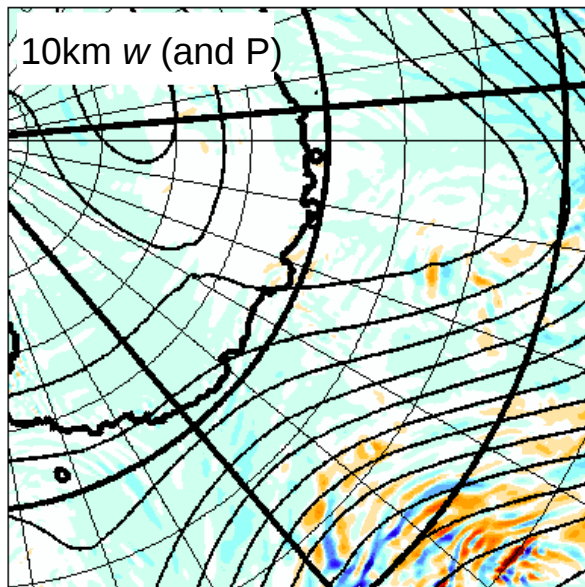
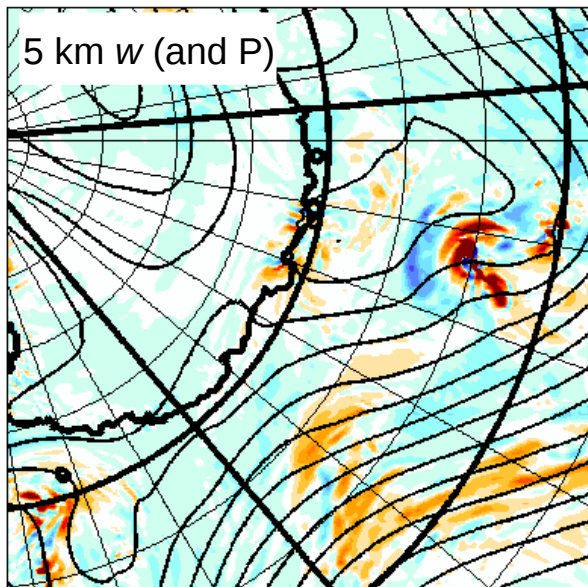
November 15, 2005, 06:00

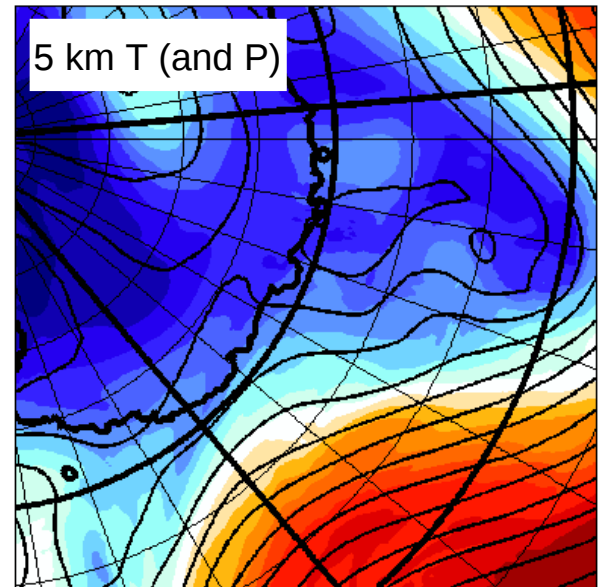
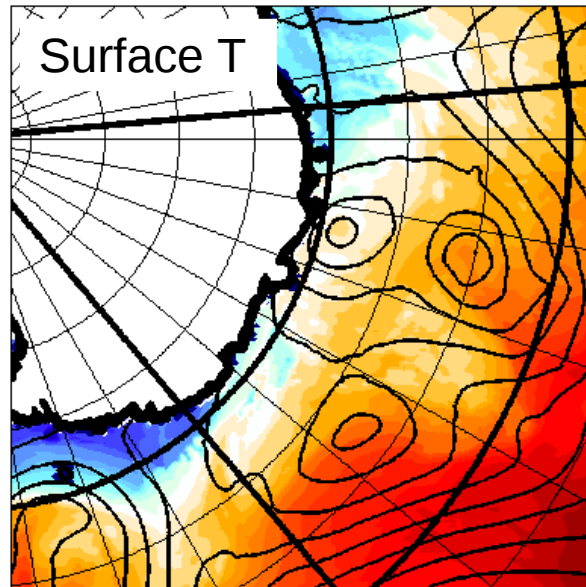
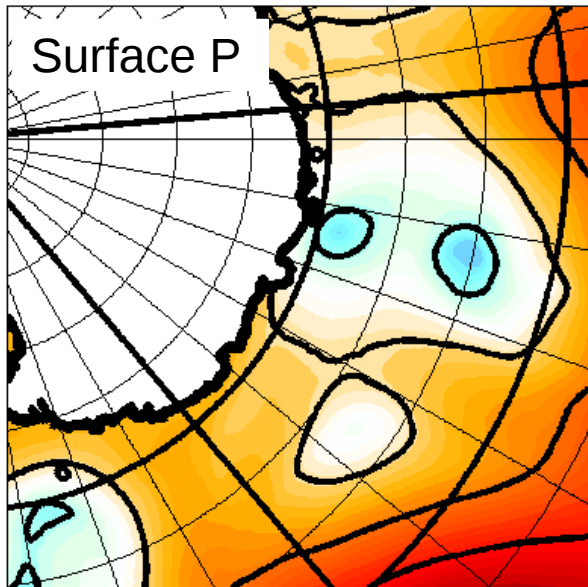




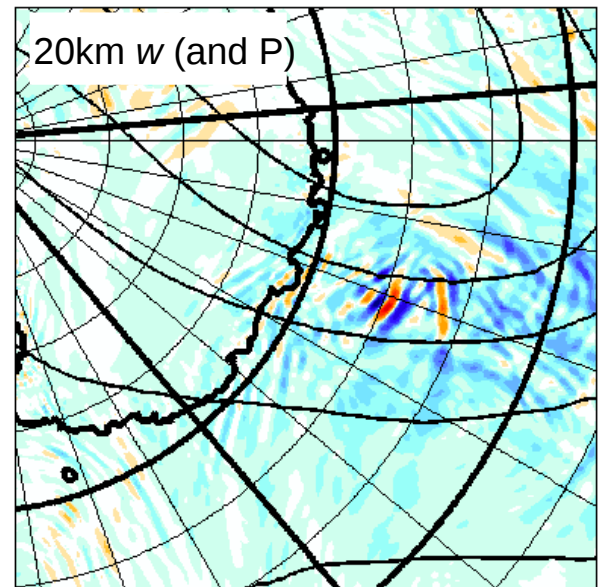
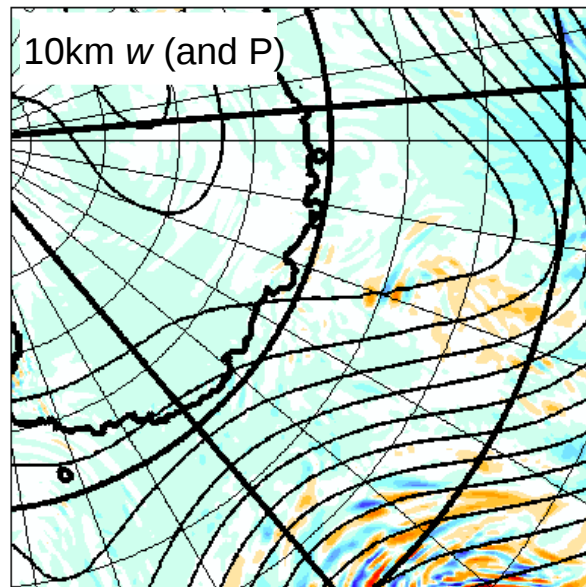
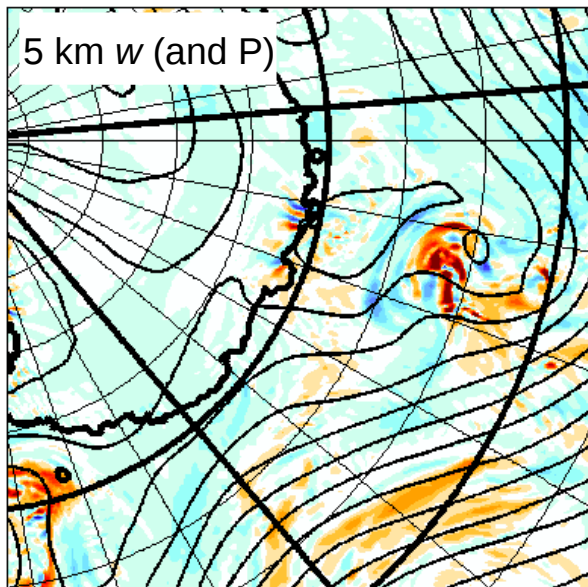


November 15, 2005, 12:00

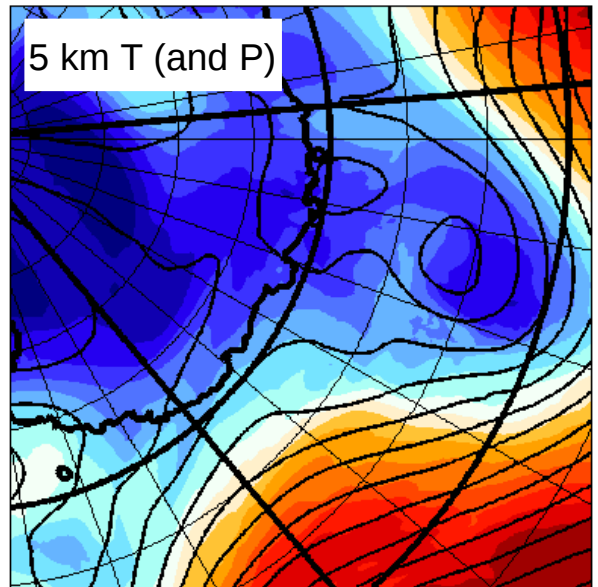
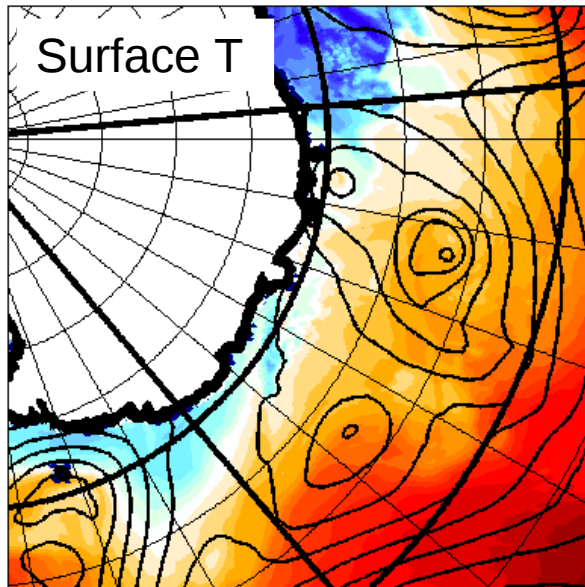
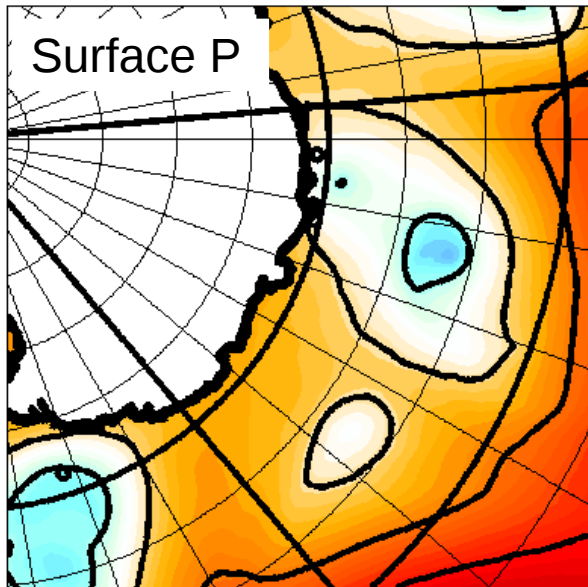




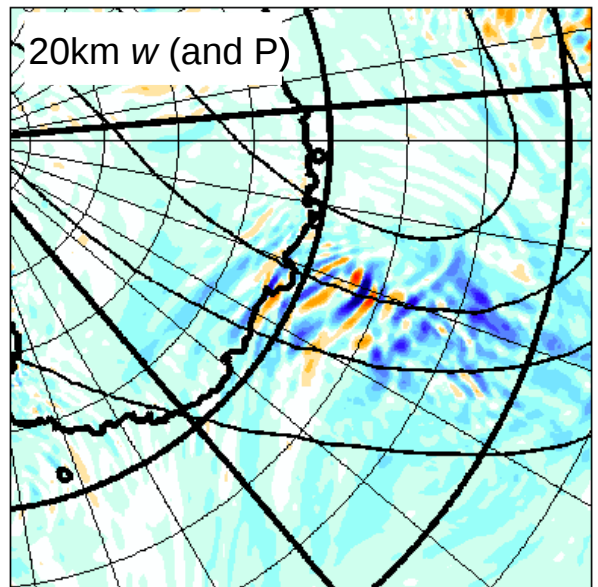
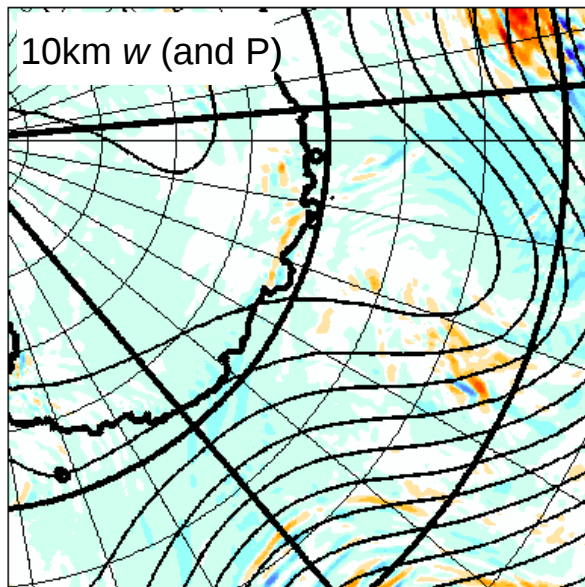
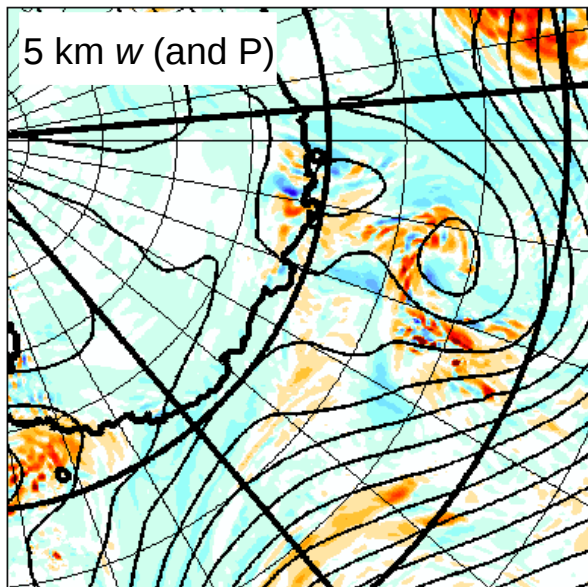
November 15, 2005, 18:00







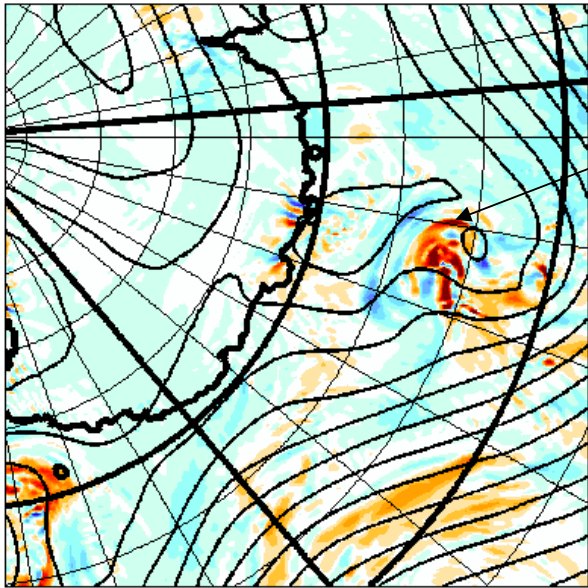
November 16, 2005, 00:00



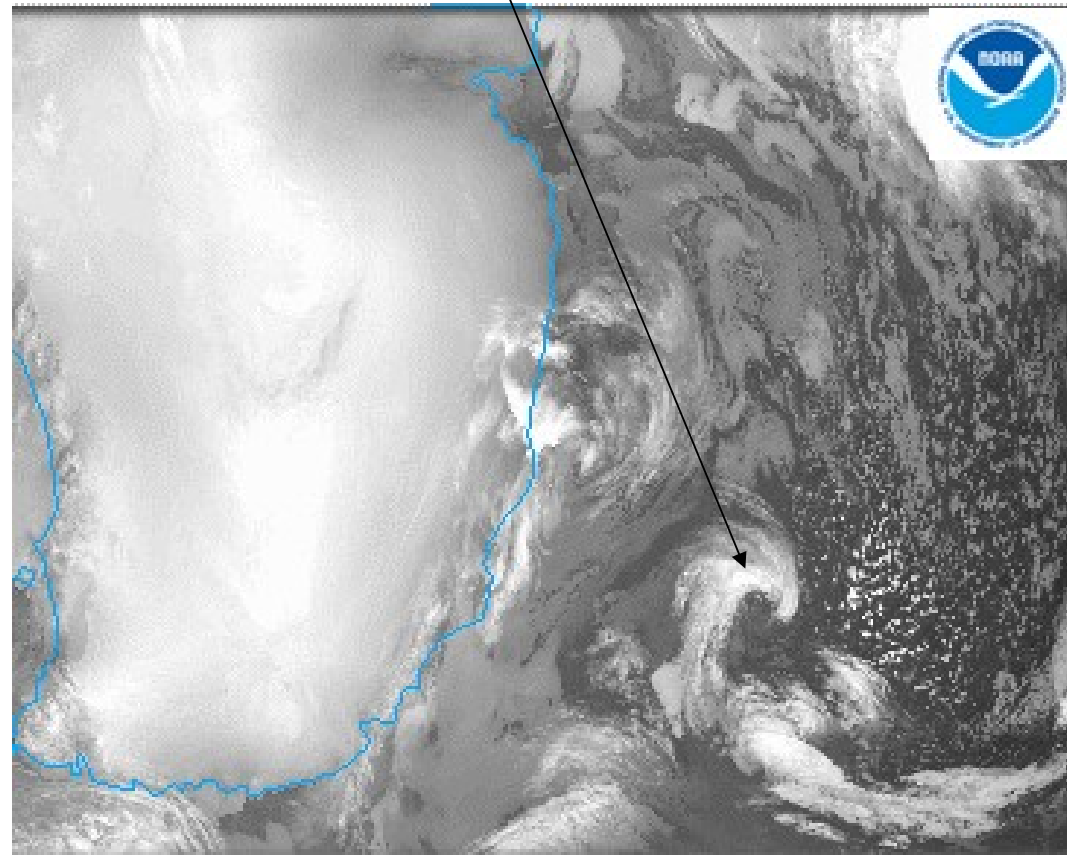
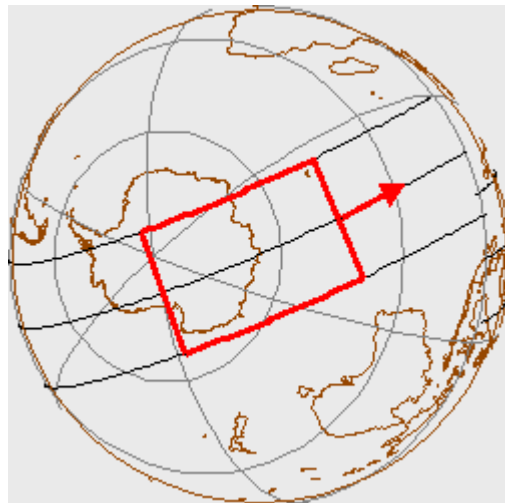


# Validation with observations : clouds around polar low

w on Nov. 15, 18UT, z=5km

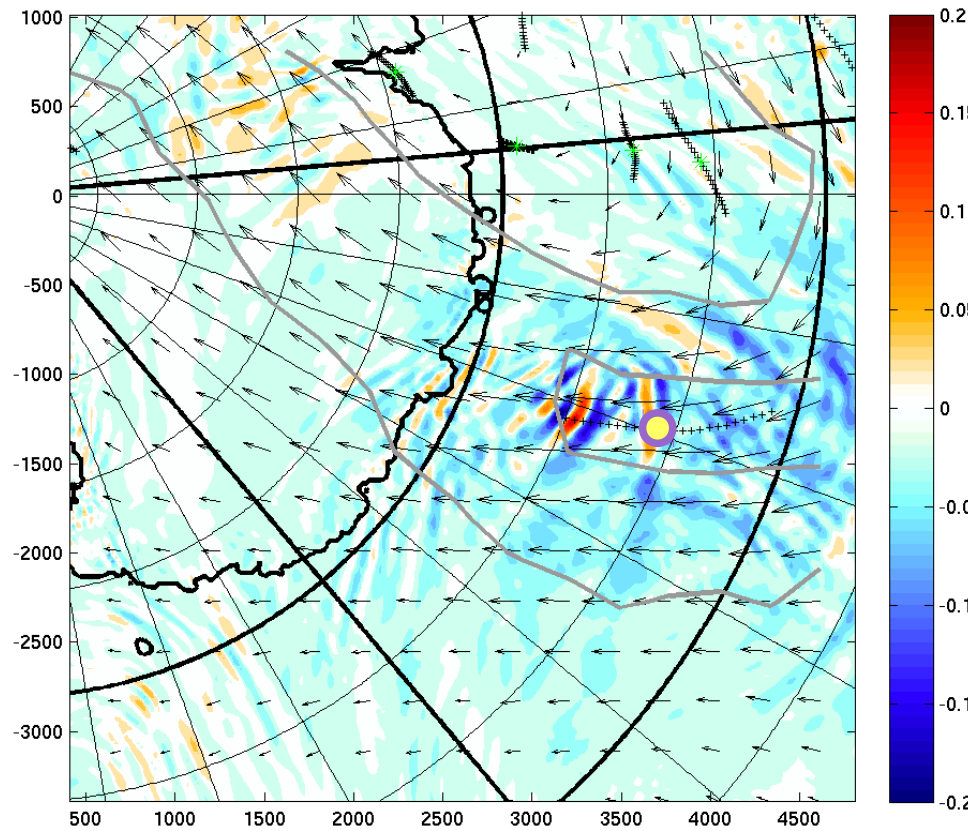


Satellite image from the Defense Meteorological Satellite Program (DMSP)



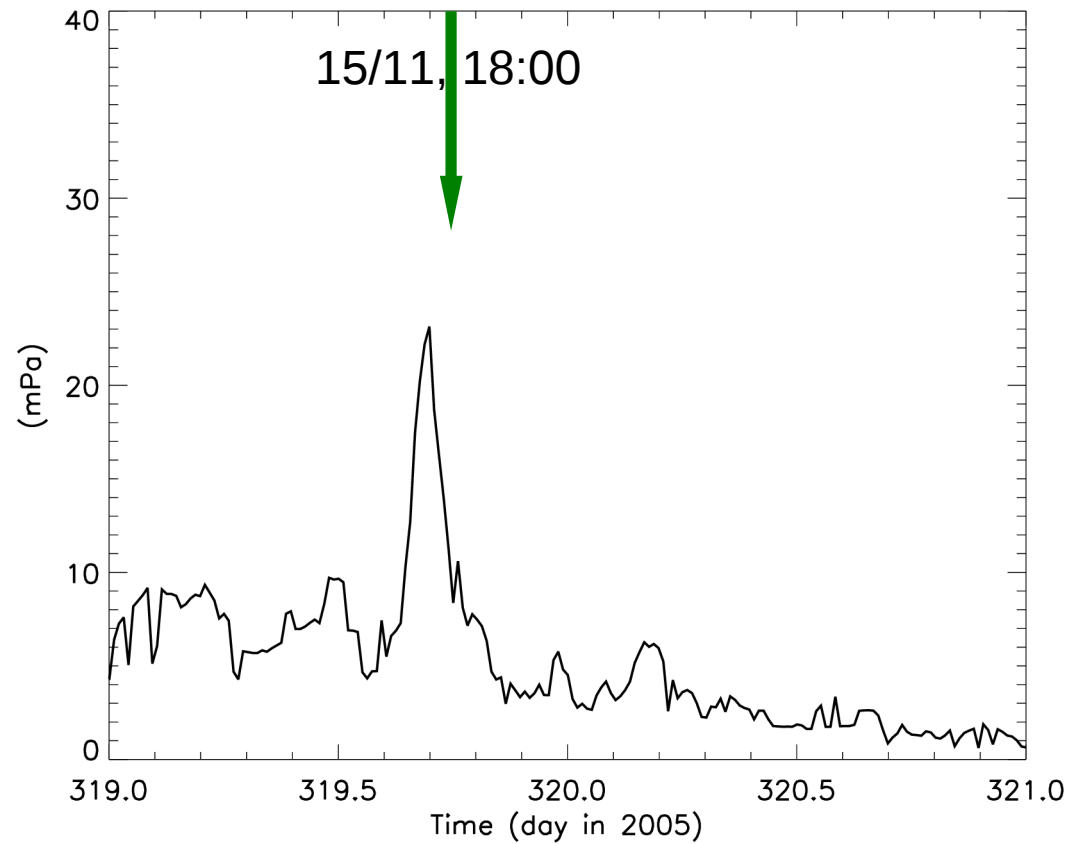
# Validation with observations : gravity waves

W (and wind) at z=20 km, 15/11, 18:00

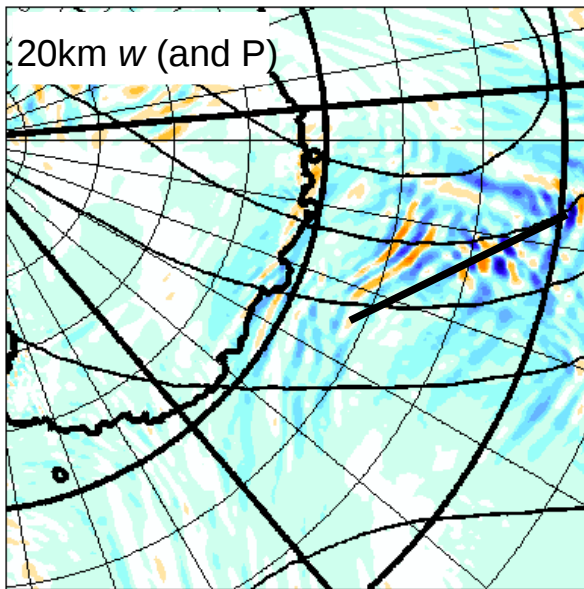


● Balloon #3

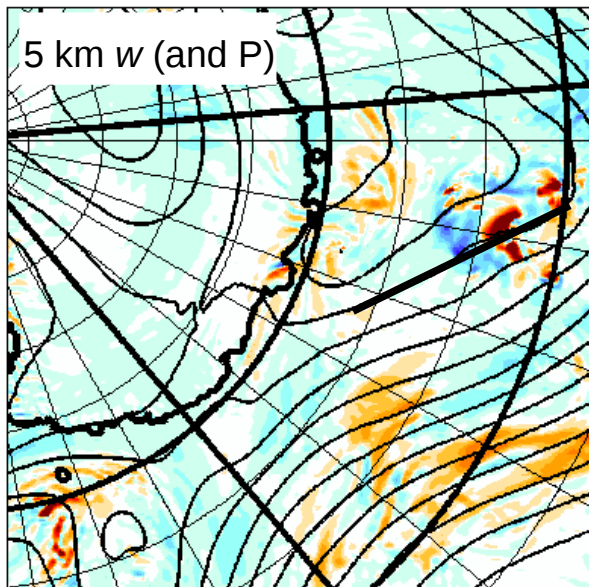
Momentum flux time series for balloon #3



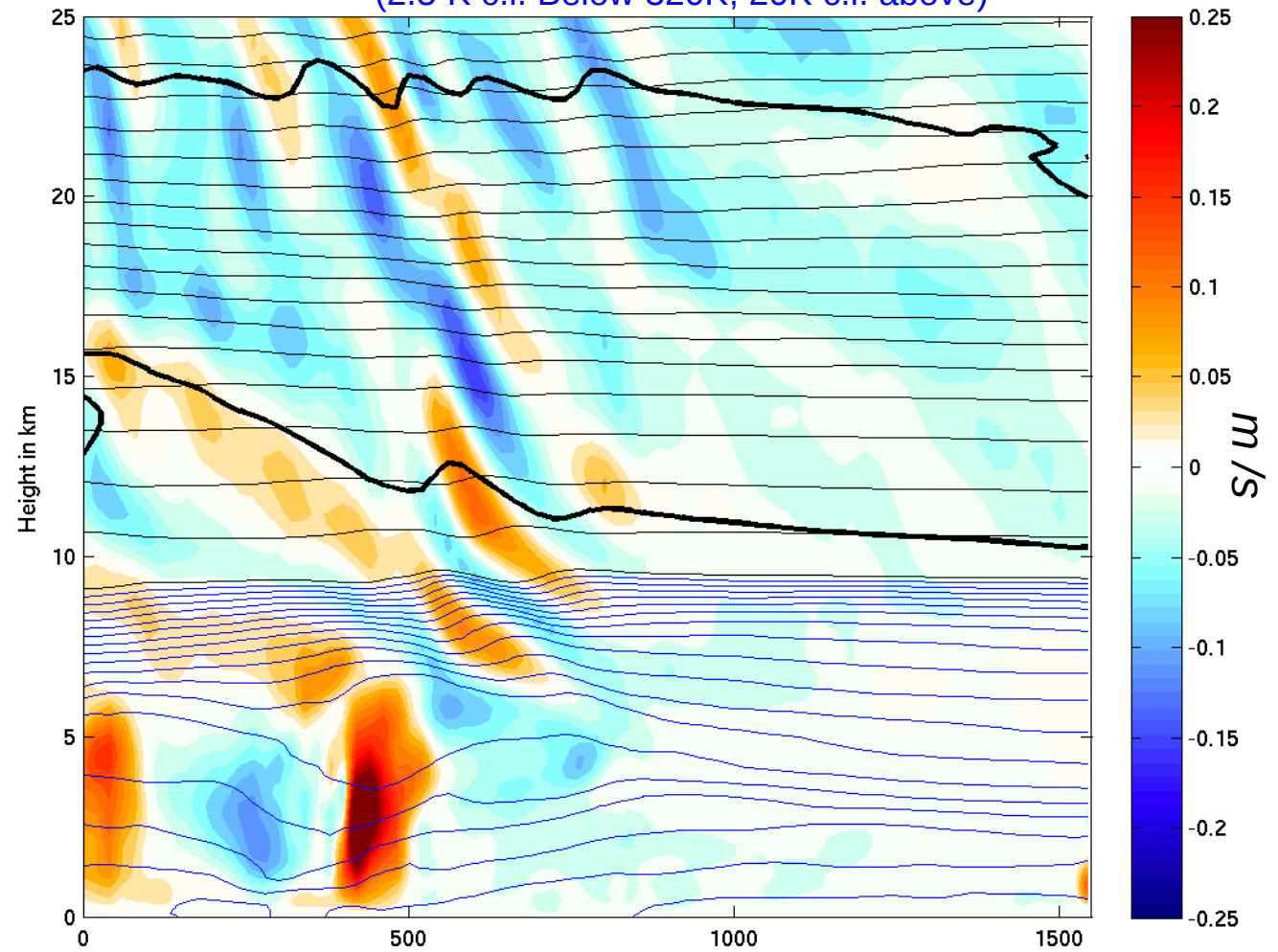
# Vertical cross-section



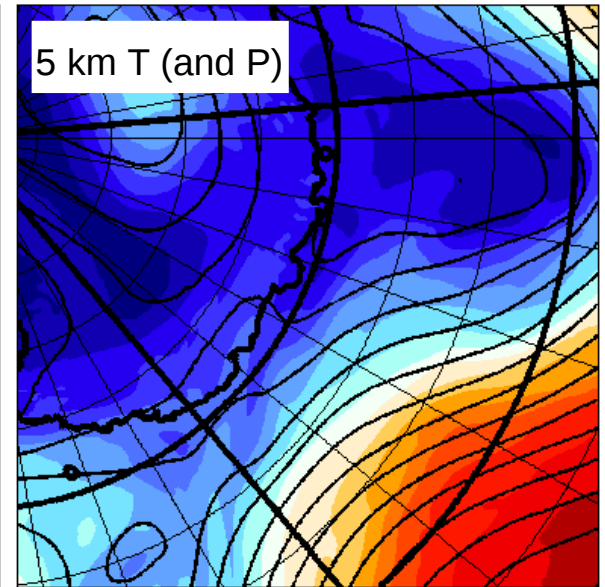
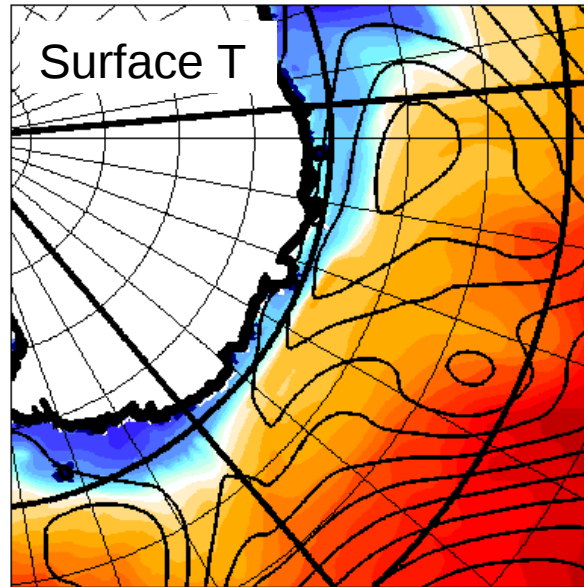
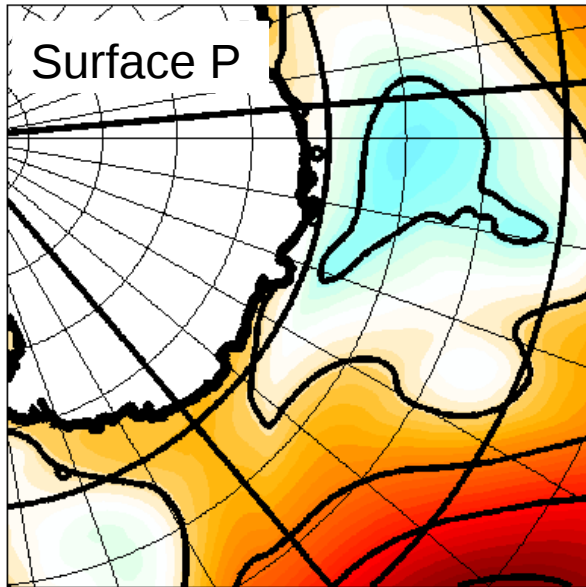
November 15, 2005, 06:00



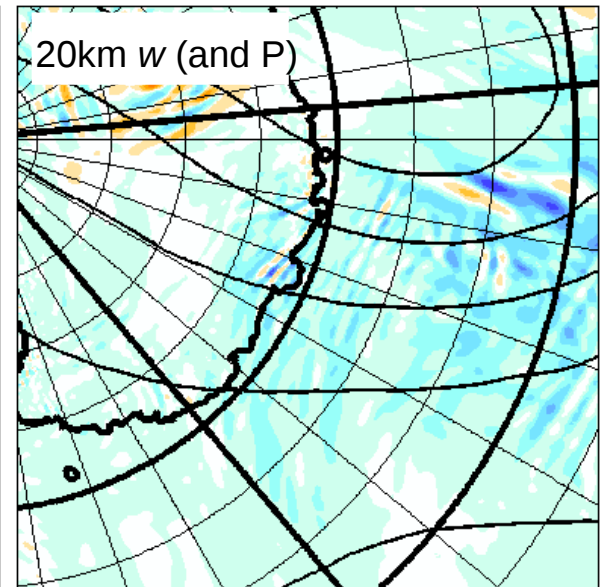
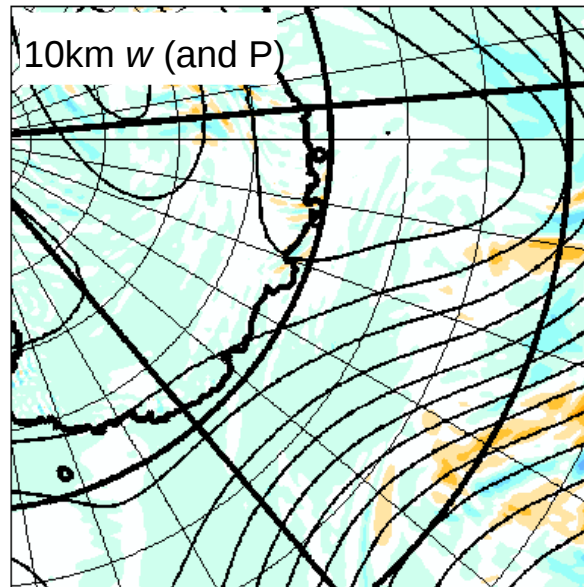
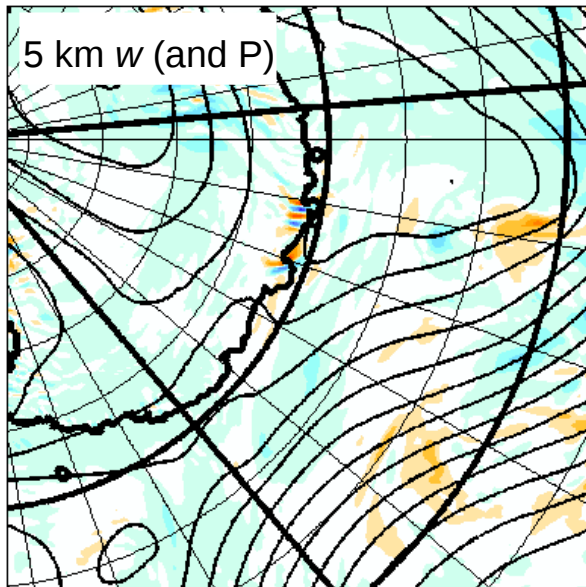
$w$ , wind speed (35 m/s isotach) and  $\theta$   
(2.5 K c.i. Below 320K, 20K c.i. above)



# Dry run for comparison



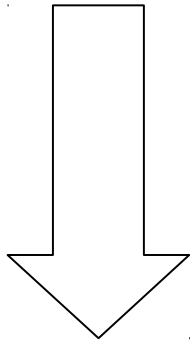
November 15, 2005, 06:00





# Comparison of dry vs moist simulations

Momentum fluxes in the dry simulations are about **2.5 times weaker**



2

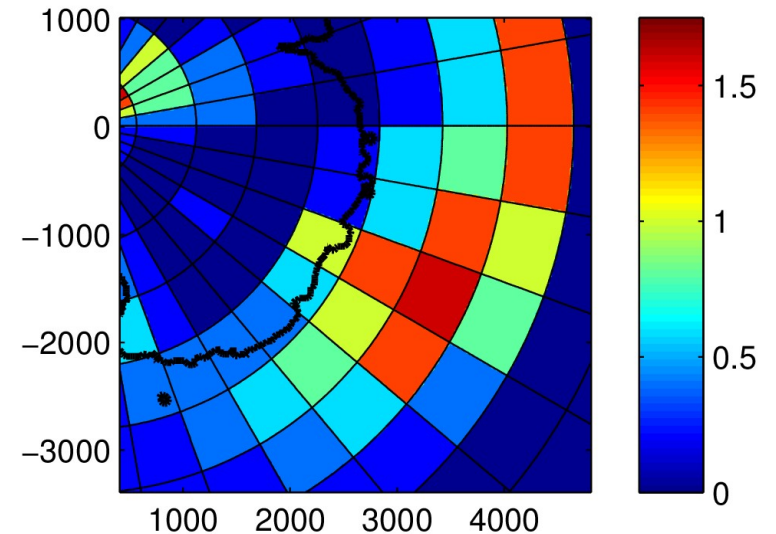
**Moist convection** near fronts appears to play a key rôle in exciting waves that are :

- **more intense,**
- **higher frequency** (5-10  $f$  rather than 1.4  $f$ )

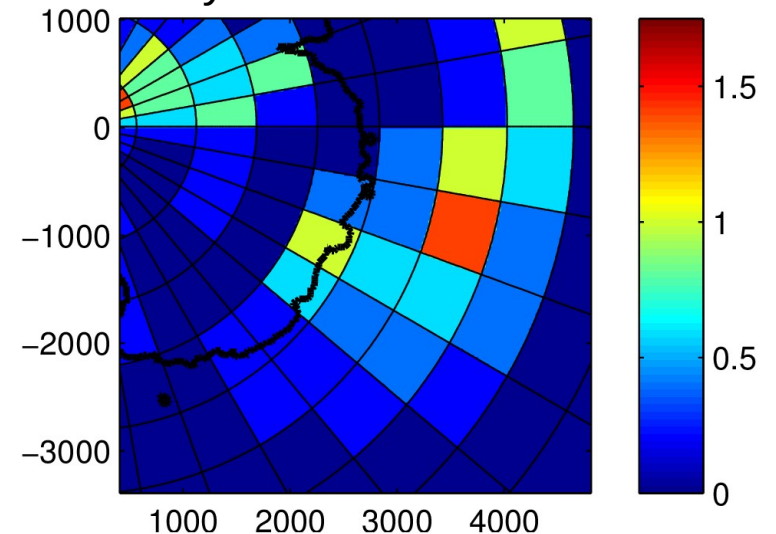
than expected from dry idealized simulations.

## 2-day mean mom. fluxes

### Full simulations

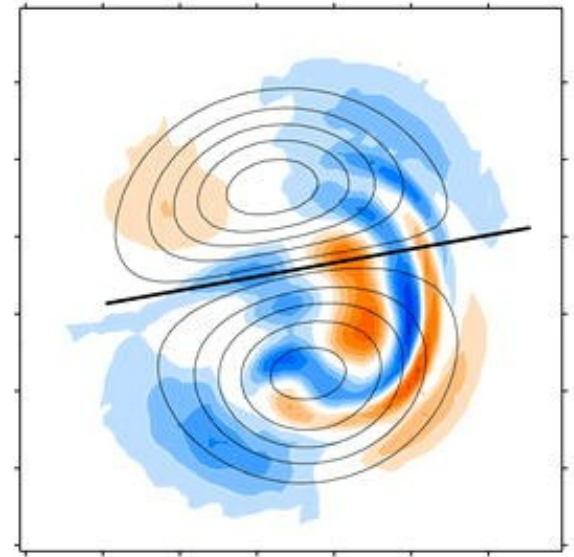
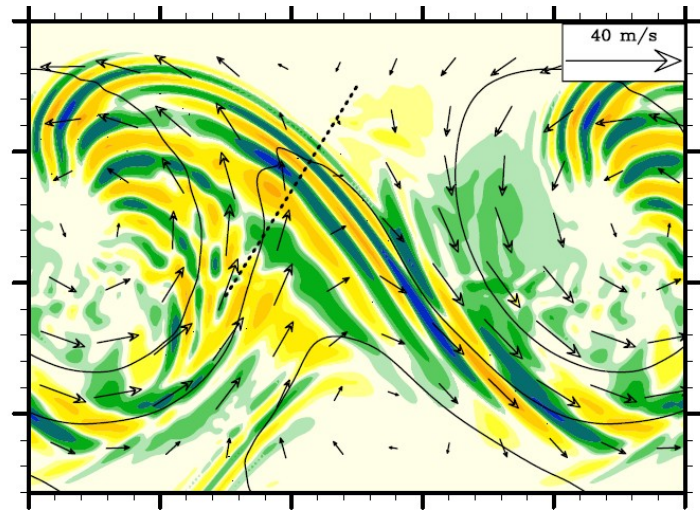
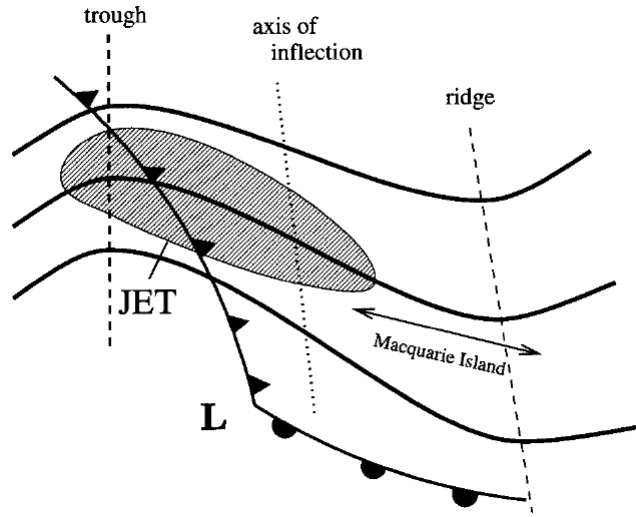


### Dry simulations





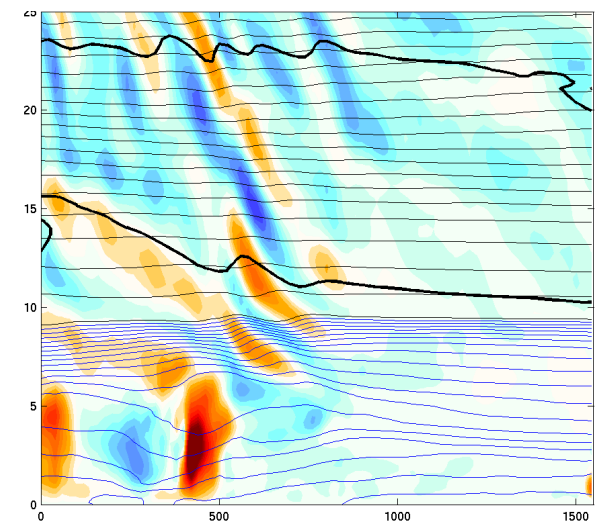
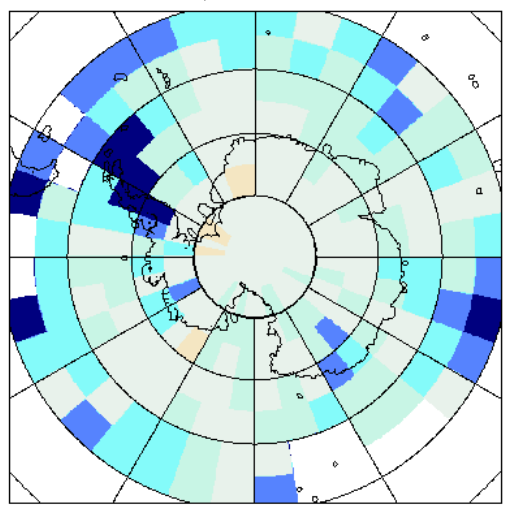
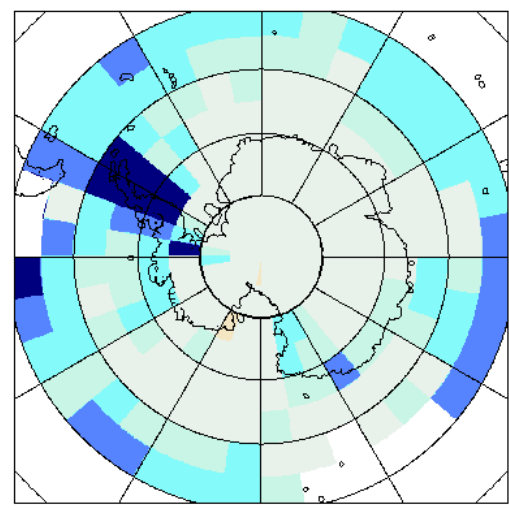




# Gravity waves generated from jets and fronts



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 Laboratoire de Météorologie Dynamique,  
 Ecole Polytechnique, Palaiseau, France



# Conclusions and Discussion

**Generation near jet exit region** well understood :

Combined roles of generation and propagation, **wave-capture**.

**Realism of simulated GW field** in mesoscale simulations

→ potential use of analyses from operational centers ?

Non-orographic GW from jets and fronts

→ **significant contribution to momentum fluxes** into the stratosphere

Case studies emphasize the **rôle of moist processes**,

→ generating **more intense, higher frequency waves**

**Remark** : criteria used to identify the waves studied :

- observations : identifiable signal, e.g. in radiosonde profiles of wind
- idealized simulations : conspicuous signal emerging in the horizontal divergence
- real case simulations : gravity wave momentum fluxes at  $z = 20$  km

Different criteria emphasize different elements of the gravity wave field ;  
jet exit region waves are also present in the real case simulations

# Conclusions and Discussion

For jet exit region waves : **backreaction of the waves on the flow**  
backreaction in the case of the dipole for example ?

**Other mechanisms** present near jets and fronts:

- generation from fronts
- small-scale shear instabilities
- rôle of **vertical shear**

Can we find a **simple relation tying GW characteristics to the large-scale flow** ?

- wave-capture : local information...
- linearized equations on background flow : costly...

**Intermittency**, and pathways for parameterizations :

- **stochastic parameterization** (*Lott et al 2012*)
- in phase with **description of the GW using PDFs** (*Hertzog et al 2012*)

What **elements of GW parameterizations are crucial to validate** ?

which simplifying assumptions are most urgent to overcome (sources, column) ?

## Thank you for your attention

A. Hertzog, M.J. Alexander and R. Plougonven (2012) On the intermittency of gravity wave momentum flux in the stratosphere, *J. Atmos. Sci.* 69 (11), p3433-3448.

R. Plougonven, A. Hertzog & L. Guez. (2013) Simulations of gravity waves above Antarctica and the Southern Ocean and comparisons to balloon observations. *Quart. J. Roy. Meteor. Soc.*, DOI:10.1002/qj.1965.

R. Plougonven & F. Zhang (2014). Internal gravity waves from atmospheric jets and fronts *Rev. Geophys*, 52, doi:10.1002/2012RG000419.

R. Plougonven, A. Hertzog and M.J. Alexander. (2015) Case studies of nonorographic gravity waves over the 1 Southern Ocean emphasize the role of moisture, *J. Geophys. Res. Atmos.*, 120, doi:10.1002/2014JD022332.